#### Primary Productivity and Nitrogen fixation by Trichodesmium 1 spp. in the Arabian Sea 2 3 Sushma G. Parab and S. G. P. Matondkar<sup>\*</sup> 4 5 National Institute of Oceanography, Dona Paula, Goa, 403 004, India. 6 Tel:. +91 (832) 2450233; Fax : +91 (832) 2450602 7 E-mail: sgpm@nio.org; psushma@nio.org 8 9 Abstract 10 Trichodesmium was studied with the purpose of understanding its distribution, organic 11 12 production and nitrogen fixation in the Arabian Sea. Out of the 143 stations sampled, a 13 total of 93 stations showed the presence of Trichodesmium filaments. Two species of 14 Trichodesmium namely, T. thiebautii and T. erythraeum were found. These were 15 distributed on the basis of the physico-chemical conditions of the Arabian Sea. This was 16 the first time that we managed to detect and record the presence of Trichodesmium 17 thiebautii bloom in the Arabian Sea at depths as deep as 70-80m. Total counts of Trichodesmium varied between 0 to 400737 filamentsL<sup>-1</sup>. Trichodesmium thiebautii 18 19 developed in offshore waters during the fall intermonsoon, when the water temperature was around $28^{\circ}$ C and nitrate content was as low as $0.34\mu$ M. After the northeast monsoon, 20 21 Trichodesmium erythraeum developed in the offshore area and then spread to coastal waters. Both species of *Trichodesmium* together produced a total of 0.263 TgCyear<sup>-1</sup> and 22 fixed a total of 0.2976 TgNyear<sup>-1</sup> in the Arabian Sea. The study revealed that 23 24 Trichodesmium was a major contributor to the organic matter productivity of the Arabian 25 Sea during the period from November to April. The seasonality of the blooms of 26 Trichodesmium is discussed with the help of Ocean Color Monitoring (OCM) data and 27 biogeochemical implication of these findings in the Arabian Sea.

28 Keywords: *Trichodesmium*; Primary production; Nitrogen fixation; Monsoon;

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#### **1. Introduction** 36

37 Trichodesmium is a planktonic filamentous cyanobacterium which is known to 38 dwell in tropical and temperate marine environments (Capone et al., 1997; Mulholland 39 and Capone 2009; Subramaniam and Carpenter, 1994). It forms dense blooms which are 40 studied with the purpose of understanding both its biomass as well as its distribution 41 (Carpenter and Prince, 1977; Capone et al., 1997). Since 1950, these blooms have been 42 reported in the Arabian Sea (Ramamurthy et al., 1972; Devassy et al., 1978; Sawant, et 43 al., 1996; Capone et al., 1998; Lugomela et al., 2002; Desa et al., 2005; Parab et al., 44 2006; Basu et al., 2011), the Lakshadweep Sea (Qasim, 1970), the Andaman Sea 45 (Jyothibabu et al., 2003) and the Bay of Bengal (Jyothibabu et al., 2003; Hegde et al., 2008). The phytoplankton biomass at times exceeds  $100\mu gL^{-1}$  during the blooming 46 period this value is higher than any other biomass value recorded in any other individual 47 48 species(Parab et al., 2006). The Trichodesmium bloom is unique since it can both synthesize organic matter and fix nitrogen (Capone et al., 1998, 2005; Carpenter et al., 49 50 1983). The ecological significance of this organism in the Arabian Sea is not very well 51 known, although it does in fact hold the key to the productivity of the Arabian Sea during 52 oligotrophic conditions.

53 The Arabian Sea has certain unique features as far as wind direction, solar insolation, currents, water temperature and upwelling during the NE (Northeast; 54 55 December to February) and SW (Southwest; June to September) monsoons are

concerned. Upwelling during the summer monsoon results in high concentrations of NO<sub>3</sub><sup>-</sup>
in surface waters, particularly in the north-western reaches of the basin. However during
the oligotrophic inter-monsoon period, surface waters are depleted of combined N
(nitrogen) thus restricting production of phytoplankton (Bhattathiri et al., 1996).

60 Bio-geochemical estimates and empirical measurements show that N<sub>2</sub> fixation by 61 Trichodesmium is a major source of new nitrogen which is of immense importance in 62 supporting the primary production of the oligotrophic environment (Capone, et al., 1997). 63 Trichodesmium spp. is considered to be an important source of new nitrogen in the 64 oligotrophic ocean; however more studies on *Trichodesmium* spp. are necessary as there 65 are large loopholes in the understanding of its seasonality in the Arabian Sea. Although this cyanobacterium is common in tropical and sub-tropical seas, its distribution is known 66 67 to be irregular, however at the same time spatially extensive dense blooms of 68 Trichodesmium form when the conditions are appropriate (Eleuterius, et al., 1981; 69 Capone, et al., 1998). Published reports (Capone, et al., 2005) on measurement of N<sub>2</sub>-70 fixation in various parts of the oceanic environment suggest that Trichodesmium is an 71 important source of new nitrogen to the phytoplankton in the ocean. Nitrogen fixation 72 studies have already been conducted and the results are available from the Atlantic 73 (Goering, et al., 1966) as well as the Pacific (Letelier, et al., 1996; Karl, et al., 1997). 74 Studies have also been conducted in the Caribbean Sea (Carpenter and Price, 1977), the 75 China Sea (Saino, et al., 1977; Chen, et al., 2003) and the Sargasso Sea (Orcutt, et al., 76 2001). It is seen that the presence of new nitrogen significantly amplifies primary 77 production and thus exports carbon into the deeper parts of the ocean. The reports on N<sub>2</sub>-78 fixation rates by Trichodesmium in the Arabian Sea are few despite the fact that,

79 Trichodesmium frequently occurs in this region (Capone et al., 1998). Also, these reports 80 have not accounted for the seasonality seen in the *Trichodesmium* blooms in the Arabian 81 Sea nor have they accounted for its diversity. Traditionally it's taken for granted that, in 82 oligotrophic conditions, the level of net biological productivity is sustained by the 83 availability of new nitrogen. This flux of new nitrogen into the euphotic zone which 84 supports primary production is balanced by the concomitant losses suffered due to 85 sinking particles, vertical migration and the mixing of organic nitrogen from the surface 86 waters (Dugdale and Goering, 1967; Eppley and Peterson, 1979; Lewis, et al., 1986; 87 Platt, et al., 1992).

88 Historical data from the Arabian Sea reveal that the cause of the low oxygen zone 89 in the Arabian Sea is partly due to upwelled water and due to respiration of organic 90 matter formed during high biological productivity. Joint Global Ocean Flux Study 91 (JGOFS) indicated high primary productivity in the Western and Central Arabian Sea 92 (Barber et al., 2001) specially during SW monsoon and NE monsoon. However Eastern Arabian Sea values are patchy ( $0.18-7.29 \text{ gCm}^{-2}\text{d}^{-1}$ ), these have been attributed to varying 93 94 nutrient (specially nitrate) and hydrographic conditions (Naqvi et al., 2006). It must also 95 be noted that the eastern Arabian Sea was not studied during the US JGOFS cruises 96 (Smith 2001, 2005). During the Indian JGOFS study primary productivity of eastern 97 Arabian Sea fluctuated seasonally. During the intermosoon, the area became oligotrophic 98 while during the winter monsoon it was found to be more productive (Bhattathiri et al., 99 1996). This area is an important site for the blooms of the *Trichodesmium* (Devassy 100 1978). The *Trichodesmium* is important as primary producers and nitrogen fixer where 101 upto 50% of the nitrogen fixed by Trichodesmium can be released as nitrogen source to phytoplankton during bloom seasons (Mulholland et al., 2006). This has prompted us to
plan study of *Trichodesmium* in the NE Arabian Sea inorder to assess seasonality and role
of the organism in the Arabian Sea ecosystem with reference to SW monsoon,
intermonsoon and NW monsoon seasons.

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#### 107 **2. Materials and Methods**

#### 108 **2.1. Study area and sampling**

109 The data presented here was collected on board two research vessels the ORV Sagar 110 Kanya (SK) and FORV Sagar Sampada (FORV). This collection of data was done in the months spanning from November to April during cruises SK-152 (30<sup>th</sup> March to 13<sup>th</sup> 111 April 2000), SK- 171 (3<sup>rd</sup> to17<sup>th</sup> November 2001); SK 186 (3<sup>rd</sup> to 19<sup>th</sup> January 2003); 112 SK214 (4<sup>th</sup> to 17<sup>th</sup> December 2004); FORV 212 (27<sup>th</sup> February to 7<sup>th</sup> March 2003) and 113 FORV 222 (22<sup>nd</sup> February to 10<sup>th</sup> March 2004). While planning the cruises, special 114 115 emphasis was given to study the distribution of Trichodesmium in the Arabian Sea covering 12-22<sup>0</sup>N and 65-74<sup>0</sup>E (Fig. 1) and seasons (November –fall inter-monsoon (fall 116 117 IntM); December to February -NE monsoon (NE MoN) and March to April - spring 118 inter-monsoon (spring IntM) so as to include samples from coastal and offshore as well 119 as regions where physical processes such as upwelling and winter cooling, were known to 120 occur. Different cruises during same seasons have been referred as the early NE monsoon 121 - December, mid NE monsoon -January, late NE monsoon-February; early spring 122 intermonsoon-March and late spring intermonsoon-April. Water samples were collected 123 from the euphotic zone with 5L Niskin sampling bottles mounted onto a Sea Bird 124 Electronics<sup>®</sup> conductivity-temperature-density profiler (CTD) rosette. The depth of the
125 euphotic zone was measured using a Secchi Disc.

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#### 129 **2.2 Physicochemical measurements**

The QUickSCAT images were downloaded from the SSM/I website at http://www.remss.com/qscat/qscat\_browse.html. Temperature and salinity data was taken from CTD. Nutrients were analyzed in an Auto-analyzer using standardized techniques for nutrient analysis in seawater (UNESCO, 1994).

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#### 135 **2.3 Chlorophyll** *a* concentration

136 To obtain the chlorophyll a (chl a) estimates seawater samples (1 litre) were 137 filtered through GF/F filters and analyzed by HPLC (Wright et al., 1991; Bidigare and 138 Charles, 2002). Prior to the HPLC analysis, the filters were immersed in 90 % acetone, 139 extracted under cold and dark conditions overnight, sonicated and finally filtered through 140 0.2 µm, 13 mm PTFE filters to free the sample from any particulate debris. Aliquots of 1 141 ml of the pigment extract were then mixed with 0.3 ml of distilled water in a 2 ml amber 142 vial and allowed to equilibrate for 5 minutes prior to its injection into an HPLC (Agilent<sup>®</sup> 1100 series) equipped with a diode array detector. Pigments were separated in a 143 144 C-18 reverse-phase column using the eluent gradient program of Wright, et al. (1991) as 145 adapted by Bidigare and Charles (2002) as detailed in Parab, et al. (2006).

#### 147 **2.4.** *Trichodesmium* spp. abundance

148 For the identification and enumeration of the phytoplankton, 500ml of seawater 149 was fixed with 1 % Lugol's iodine and preserved in 3 % buffered formaldehyde solution 150 and then stored under cool and dark conditions until the time of analysis. Prior to 151 microscopic analysis, the samples were concentrated to about 5- 10 ml by carefully 152 siphoning the top layer with a tube covered by a 10 µm Nitex filter on one end. 153 Replicates of 1ml sample concentrates were transferred to a Sedgwick-Rafter slide and 154 counted using an Olympus Inverted microscope (Model IX 50) at 200X magnification. 155 The smallest cells enumerated were of ~5µm in diameter. Identifications were based on 156 standard taxonomic keys (Tomas et al., 1997). The results were expressed as the number 157 of cells with the exception of *Trichodesmium* which was expressed as the number of 158 trichomes/filaments.

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#### 160 **2.5. Primary productivity**

161 For primary productivity analysis, water samples were collected depending upon 162 the euphotic zone that was established using a Secchi Disc. For total phytoplankton 163 measurement, the volume of water samples was transferred in triplicates to 250 ml capacity Nalgene Polycarbonate bottles. One ampoule of <sup>14</sup>C (5µci) was injected to each 164 165 three light and one dark bottles. These bottles were then incubated for 12 hrs on deck 166 using appropriate light cutoff screens. At the end of the incubation period the water 167 sample (250 ml) was filtered through GF/F filter. All the filters exposed to fuming HCL 168 for 20 minutes in a desiccator inorder to dissolve left over bicarbonates. Samples were 169 then placed in a 7 ml mini scintillation vial with scintillation cocktail and activity was determined with a Wallac Scintillation counter (model no. 1409). All counts were quenchcorrected (UNESCO, 1994).

To measure *Trichodesmium* productivity, trichomes were first collected in a 202 µm mesh 1 m diameter net towed. Trichomes/filaments were picked up and rinsed in filtered seawater. The known numbers of trichomes were then placed in a Nalgene Polycarbonate bottles containing 250 ml of filtered seawater from same station and <sup>14</sup>C (5µci) was added. Further, the same procedure was followed as described for phytoplankton productivity. The phytoplankton and *Trichodesmium*, daily carbon fixation was measured using standard (UNESCO, 1994).

#### 179 **2.6.** Nitrogen fixation

180 *Trichodesmium* spp. filaments counted by inverted microscope were further used 181 to calculate the rate of nitrogen fixed per filament using factor developed by Capone et 182 al., 1998 for the Arabian Sea and further validated by C/N content of filaments (for fixed 183 period) using CHN analyzer. Nitrogen fixed was calculated per 10 hrs and then per year.

- 184 **2.7. Satellite studies**
- 185 The Chl *a* images were generated using OC-2 algorithm; ERDAS Imagine 2.5186 software was used for image processing.

# 187 **3. Results**

- 188 **3.1 Fall Intermonsoon**
- 189 3.1.1. Wind, temperature, salinity and nutrients

During the fall inter-monsoon the wind was blowing from the northeast towards the offshore region of the Arabian Sea and wind speed was found to be 5-6 ms<sup>-1</sup> (Fig. 2a). These calm conditions supported the dense bloom of the *Trichodesmium thiebautii* at 17<sup>0</sup>

193	N $70^{0}$ E and $17^{0}$ N $69^{0}$ E. In this area the water temperature was found to be $28^{0}$ C while
194	mixed layer depth was 5m (coastal) to 70 m (offshore) region (Fig. 3a).
195	The area of Trichodesmium thiebautii was marked by occurrence of nitrate and
196	nitrite at non detectable level (Figs. 4a, 4b & 5a). The phosphate and silicate level at
197	these areas was also low, registering at $\sim 0.2~\mu M$ and $\sim 0.5~\mu M$ respectively. Other areas
198	$(15^{0}N 73^{0}E \text{ and } 18^{0}N 69^{0}E)$ showed high level of nutrients with high percentage of
199	diatoms and absence of Trichodesmium.
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205	Insert Fig. 4 here
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207	Insert Fig. 5 here
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210	3.1.2. Chlorophyll a and Trichodesmium distribution
211	Chl <i>a</i> content was high $(0.1733 \text{ to } 5.344 \text{ mgm}^{-3}; \text{ Fig. 6a})$ at most of the locations
212	during the fall IntM. The surface Chl $a$ maxima observed (Fig. 7a) at areas like $18^{\circ}$ N
213	$70^{0}E$ (0.75 mgm <sup>-3</sup> ) and $20^{0}N$ $71^{0}E$ (5.344 mgm <sup>-3</sup> ) was on account of the blooms of
214	<i>Trichodesmium thiebautii</i> in the offshore region. Similarly at $17^{0}N 69^{0}E$ and $18^{0}N 70^{0}E$
215	up to a depth of 30 m, Trichodesmium bloom was noticed. Out of the 23 stations
216	sampled 10 stations showed the presence of Trichodesmium thiebautii. The highest

217	phytoplankton counts were observed at $18^{0}$ N $70^{0}$ E (15.831 cell nos. x $10^{4}$ L <sup>-1</sup> ) where
218	Trichodesmium constituted a major portion. The average total cells counts of
219	phytoplankton was $1.590\pm3.546 \times 10^4 L^{-1}$ at the surface and $43.226\pm95.189 \times 10^6 m^{-2}$ in the
220	euphotic zone (Fig. 6b). The number of filaments across transect varied from 0 to
221	10.9570 filaments x $10^4$ L <sup>-1</sup> . The average filament distribution was 6.9993±3.025 x $10^4$ L <sup>-1</sup>
222	at the surface and 9.686 $\pm$ 30.720 x 10 <sup>6</sup> m <sup>-2</sup> in the euphotic zone (Fig. 6c). At some areas
223	like, $16^{0}$ N $71^{0}$ E, $19^{0}$ N $69^{0}$ E and $20^{0}$ N $70^{0}$ E N the diatom count reached right up to 100 %.
224	However, the Trichodesmium stations were invariably marked by low percentages of
225	diatoms and dinoflagellates.

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228 Insert Fig. 7 here

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# 230 3.1.3. Primary Production and Nitrogen Fixation

231 Primary productivity measured during the fall IntM is presented in Fig. 8a. The surface maxima of high primary productivity at  $20^{\circ}$ N  $70^{\circ}$ E and  $21^{\circ}$ N  $68^{\circ}$ E was due to the 232 *Trichodesmium* population. Although, at  $21^{\circ}$ .1N the concentration of Chl *a* was high at 233 234 40 m due to the presence of the Trichodesmium bloom productivity was found to be 235 negligible due to low light at the bottom of the euphotic zone. At other stations, it was 236 observed that the sub-surface Chl a maxima due to other phytoplankton rarely coincided 237 with primary productivity maxima. The carbon fixed by Trichodesmium varied from 0.43 to 20.487 mgCm<sup>-3</sup>d<sup>-1</sup>. The rate of carbon production was high at 18<sup>0</sup>03.9N where 238 temperature was  $29^{\circ}$ C. The average carbon produced during this season was  $8.460\pm$ 239

6.504 mgCm<sup>-3</sup>d<sup>-1</sup>. During our study 12.8 % of the carbon was produced by *Trichodesmium* and the remaining 87.2 % was produced by other phytoplankton as primary producers. The highest value of primary productivity (20.487 mgCm<sup>-3</sup>d<sup>-1</sup>) was recorded at the surface of  $18^{0}$ N  $70^{0}$ E, where active growth of *Trichodesmium thiebautii* taking place.

The N<sub>2</sub> fixed during this study was  $114.110 \pm 345.770 \ \mu mol \ Nm^{-3}d^{-1}$  and varied from 0 to 1252.390  $\mu$ molNm<sup>-3</sup>d<sup>-1</sup> (Fig. 8b). Higher amounts of nitrogen was fixed at  $18^{0}$ N 70<sup>0</sup>E due to the *Trichodesmium* bloom.

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## **3.2 NE Monsoon**

#### 251 3.2.1. Wind, temperature and salinity

252 During the NE Monsoon the wind was blowing from the northswestt and moving 253 along the eastern coast of India (Fig. 2b). The wind speed during this period was found to be <5ms<sup>-1</sup>. During this season offshore water were showing the effects of winter cooling 254 255 and it was found that the growth of the Trichodesmium thiebautii continued at this time. 256 The temperature data collected, during NE MoN from the NE Arabian Sea concerning 257 temperature is presented in Fig. 3b. The observations and analysis of data made in the 258 present study indicate that the offshore water was largely influenced by the winter cooling lowering the temperatures (Fig. 3b). During this season, Trichodesmium 259 260 thiebautii began to appear in the shelf region where the impact of winter cooling had 261 diminished. By the end of the NE MoN it was found that Trichodesmium erythraeum had 262 assumed the role of the dominating species of the shelf region.

Nutrients concentrations recorded during this season was low (Figs. 4 & 5). However, Trichodesmium erythraeum was recorded alongside of Trichodesmium 264 265 thiebautii specially in the shelf region. The prevailing calm weather during this period 266 was particularly responsible for the growth of *Trichodesmium* at the sub-surface. The 267 phosphate varied from 0 to 0.35  $\mu$ M which was low most of the time. The nitrate values 268 ranged from 0 to 0.75 µM and the silicate content recorded was 11.08 µM which was 269 high at the bottom of the euphotic zone, as compared to the surface which was  $2.30 \mu M$  ( 270 Figs. 4d and 5).

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#### 272 3.2.2. Chlorophyll a and Trichodesmium distribution

273 Phytoplankton biomass in terms of the Chl *a* and total cell counts was reasonably high during this season (Fig. 6a). Average Chl *a* was  $0.526\pm0.453$  mgm<sup>-3</sup> at surface and 274  $32.588 \pm 46.851 \text{ mgm}^{-2}$  in euphotic zone (Fig. 6a). Highest Chl *a* was recorded at  $21^{\circ}$ N 275  $69^{0}$ E and lowest at  $12^{0}$ N  $72^{0}$ E. Similarly average cell density at the surface was 276  $0.112\pm0.129 \text{ nos.} x10^4 \text{L}^{-1}$  (Fig. 6b). The highest cell counts (0.606 nos.  $x10^4 \text{L}^{-1}$ ) were 277 recorded at  $12^{0}N 74^{0}E$  and lowest (0.002 nos.x $10^{4}L^{-1}$ ) at  $21^{0}N 68^{0}E$ . Filament counts 278 of Trichodesmium spp. varied from 0 to 0.616 filaments nos.x10<sup>4</sup>L<sup>-1</sup> and average 279  $0.039\pm0.119$  filaments nos.x10<sup>4</sup>L<sup>-1</sup> at the surface and  $0.214\pm0.272 \times 10^{6} \text{m}^{-2}$  at the euphotic 280 zone (Fig. 6c). 281

282 A district seasonality was recorded during SK-214, SK-186, FORV-212 and 283 FORV-222 undertaken. During the early NE MoN (SK-214), the concentration of Chl a was from  $0.187 - 0.507 \text{ mgm}^{-3}$ . High concentration of Chl *a* was recorded at  $17^{\circ}\text{N}$  69°E 284 while low concentration of Chl *a* was found at  $15^{\circ}$ N  $71^{\circ}$ E. The sub-surface Chl *a* 285 maxima was found at a depth of 40 m depth at  $17^{\circ}N$   $69^{\circ}E$ ,  $18^{\circ}N$   $67^{\circ}E$ ,  $20^{\circ}N$   $66^{\circ}E$ ,  $21^{\circ}N$ 286

69<sup>°</sup>E and 21<sup>°</sup>N 65<sup>°</sup>E (Fig. 7b). Out of the 14 stations sampled, 10 samples showed the 287 288 presence of *Trichodesmium* in the sub-surface waters while others showed its presence at 289 deeper depths (~ 60 m). The total phytoplankton cell counts varied from 0.032 to 0.4352 cell nos x  $10^4 L^{-1}$  at the surface. At two locations ( $15^0 N 73^0 E$  and  $18^0 N 70^0 E$ ) 290 Trichodesmium erythraeum and Trichodesmium thiebautii were found at the depth of 5 291 292 and 60 m respectively. The highest counts of Trichodesmium filaments were found at  $22^{0}$ N 67<sup>0</sup>E (0.0216 filaments x 10<sup>4</sup>L<sup>-1</sup>) and lowest at 20<sup>0</sup>N 66<sup>0</sup>E (0.0022 filaments x 10<sup>4</sup>L<sup>-1</sup>) 293 <sup>1</sup>). During this season diatoms were also recorded in large numbers at  $15^{\circ}N$   $73^{\circ}E$ ,  $15^{\circ}N$ 294 71°E and 18°N 67°E. The Trichodesmium percentages in total phytoplankton counts 295 296 ranged from 0.78 to 35.71 %. The other dominant phytoplankton throughout this season 297 were Rhizosolenia setigera, Nitzschia longissima, Thalassiothrix longissima, 298 Pyrophaccus horologium and Protoperidinium sternii (Table 1).

299 The Chl *a* during the mid NE MoN (SK-186) was much higher at most of stations, compared to early NE MoN. It ranged from 0.173 to 2.6mgm<sup>-3</sup>. Higher concentrations of 300 chl *a* were found at  $21^{\circ}$ N 69°E where diatoms were dominant (93.94 %) followed by 301 302 dinoflagellates (4.55 %) and other algae like (1.52 %). Sub-surface Chl a maxima was observed (Fig. 7b) at some locations 17<sup>0</sup>N 72<sup>0</sup>E, 17<sup>0</sup>N 70<sup>0</sup>E, 17<sup>0</sup>N 67<sup>0</sup>E, 19<sup>0</sup>N 66<sup>0</sup>E, 20<sup>0</sup>N 303 66°E, 20°N 66°E, and 17°N 68°E during this season. The total phytoplankton cell counts 304 ranged from 0.002 to 0.459 cell nos x  $10^4 L^{-1}$ . Higher counts were found at  $17^0 N 70^0 E$ 305 306 where, diatoms made up 94.44 % and dinoflagellates formed 5.56% of the total 307 phytoplankton population. The presence of Trichodesmium (25 to 28.57 %) was found in 308 the 5 out of the 25 stations sampled at the surface. Trichodesmium erythraeum was found 309 to be more dominant than Trichodesmium thiebautii. The Trichodesmium filaments was more (760 filaments L<sup>-1</sup>) at 18°N 69°E. The other species that were observed during this
period were *Thalassiothrix longissima*, *Chaetoceros dichaetea*, *Thalassiosira subtilis*, *Coscinodiscus nitidus*, *Navicula* spp., *Nitzschia seriata*, *Nitzschia closterium*, *Rhizosolenia setigera* and *Chaetoceros lorenzianus* (Table 1).

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#### 315 Insert Table 1 here

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During late the NE MoN (part of FORV-212 and FORV-222) it was found that 317 values for Chl *a* varied from  $0.132 - 0.960 \text{ mgm}^{-3}$  with high values at  $17^{0}\text{N} 69^{0}\text{E}$  due to 318 319 the presence of the Trichodesmium erythraeum bloom. The sub-surface Chl a maxima was found at a depth of ~ 40-60 m specially at  $16^{\circ}$ N  $72^{\circ}$ E,  $10^{\circ}$ N  $74^{\circ}$ E,  $12^{\circ}$ N  $72^{\circ}$ E and 320 14°N 73°E. The presence of *Trichodesmium* was observed in 7 out of the 10 stations 321 322 sampled in this month. The total phytoplankton counts ranged from 0.0078 to 0.2508 cell nosx10<sup>4</sup>L<sup>-1</sup>. Higher phytoplankton cell counts were found at 17<sup>0</sup>N 69<sup>0</sup>E due to the 323 presence of the Trichodesmium erythraeum bloom. The Trichodesmium filaments ranged 324 from 0 to 0.6160 filaments x  $10^4 L^{-1}$ . Further a *Noctiluca* bloom was also recorded with 325 the *Trichodesmium* bloom at  $10^{\circ}$ N  $64^{\circ}$ E and  $17^{\circ}$ N  $69^{\circ}$ E. It appears that the *Noctiluca* 326 327 bloom may have followed the Trichodesmium bloom in the Arabian Sea, an idea which 328 provides us with new insights in the study of phytoplankton in the Arabian Sea. A full growth of *Noctiluca miliaris* bloom was stuck at  $17^{0}$ N  $70^{0}$ E. 329

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# 334 3.2.3. Primary Productivity and Nitrogen Fixation

The primary productivity during this period was high and ranged from 0.907-2572.1 mgCm<sup>-3</sup>d<sup>-1</sup> (av. 217 mgCm<sup>-3</sup>d<sup>-1</sup>) at surface and 4.284 mgCm<sup>-2</sup>d<sup>-1</sup> in the euphotic zone (Fig. 8a).

The primary productivity maxima at  $18^{\circ}$ N  $70^{\circ}$ E (40 m) coincided with the 338 *Trichodesmium* bloom in the water column. Similarly, at 19<sup>0</sup>N 70<sup>0</sup>E, high productivity at 339 340 both, the surface and below 40 m was also due to Trichodesmium. The average rate of 341 primary production during the bloom of *Trichodesmium* spp. was  $0.356 \pm 0.096$  mgCm<sup>-</sup>  $^{3}d^{-1}$ . At the surface, the carbon production varied from 55.85 to 706.66 mgCm<sup>-3</sup>d<sup>-1</sup>. 342 Higher rate of primary production was observed at 22<sup>0</sup>N 67<sup>0</sup>E, due to the dense bloom of 343 the Trichodesmium. The primary production by Trichodesmium thiebautii and 344 345 Trichodesmium erythraeum during the NE monsoon was of the same order i.e. 0.392 and  $0.378 \text{ mgCm}^{-3}\text{d}^{-1}$  respectively. When *T. erythraeum* bloom was prominent in the water 346 column however it was found that primary production was high  $(10.059 \text{ mgCm}^{-2}\text{d}^{-1})$ 347 compared to the primary production by T. thiebautii (6.156 mgCm<sup>-2</sup>d<sup>-1</sup>). The percentage 348 349 of primary productivity due to Trichodesmium spp. in the total productivity was around 350 0.11 % .

During mid NE MoN (SK-186), primary productivity was low at the surface and high at deeper depths. At  $15^{0}$ N  $72^{0}$ E, the highest primary productivity and Chl *a* was observed at a depth of 50 m due to the *Trichodesmium* bloom. Similarly carbon production was observed at  $18^{0}$ N  $69^{0}$ E at 60 m due to the bloom of the *Trichodesmium*  355 *thiebautii*. It was observed that *Trichodesmium* played a role of primary producer at the 356 surface like at  $22^{0}$ N  $68^{0}$ E as well as throughout the water column.

The amount of primary production due to *Trichodesmium* during late NE MoN varied from 1.61 to 8.25 mgCm<sup>-3</sup>d<sup>-1</sup>. The average amount of carbon produced was 7 mgCm<sup>-3</sup>d<sup>-1</sup>. The observations made during this season indicate that *Trichodesmium* contributed a major portion (60 %) to the carbon produced at this time. Higher amount of carbon production (65.85 %) by *Trichodesmium* was seen at area  $14^{0}N$  73<sup>0</sup>E while lower production was recorded at  $13^{0}N$  70<sup>0</sup>E (0.14 %).

The N<sub>2</sub> fixation during this season ranged from 0 -13.598  $\mu$ molNm<sup>-3</sup>d<sup>-1</sup> at the 363 surface and 0 - 31.597  $\mu$ molNm<sup>-2</sup>d<sup>-1</sup> at the euphotic zone and depended upon the 364 intensity of the bloom. The average  $N_2$  fixed during this season was  $4.432\pm13.598$ 365  $\mu$ molNm<sup>-3</sup>d<sup>-1</sup> (Fig. 8b). The rate of N<sub>2</sub> fixation also varied species wise. The amount of 366 nitrogen fixed by T. erythraeum varied from 0 to 2.4  $\mu$ mol Nm<sup>-3</sup>d<sup>-1</sup>. In the column, the 367 nitrogen fixed by T. erythraeum was found to be much higher  $(23 \pm 4.77 \text{ }\mu\text{mol }\text{Nm}^{-2}\text{d}^{-1})$ 368 than that fixed by T. thiebautii (  $14.09 \pm 4.074 \mu mol Nm^{-2}d^{-1}$ ). During the late NE MoN 369 the total nitrogen fixed by Trichodesmium spp. varied from 0 to 70.409 µmolNm<sup>-3</sup>d<sup>-1</sup> 370 with an average of  $14.002 \pm 23.772 \text{ } \mu\text{molNm}^{-3}\text{d}^{-1}$  (n=8). Higher amount of nitrogen was 371 fixed at  $12^{0}$ N  $74^{0}$ E and lower amount of it was fixed at  $16^{0}$ N  $72^{0}$ E. The amount of N<sub>2</sub> 372 fixed by *T. erythraeum* varied from 1.25 to 37.25 µmolNm<sup>-3</sup>d<sup>-1</sup> which was much higher 373 than that fixed by *T. thiebautii* (0 to 32.67  $\mu$ molNm<sup>-3</sup>d<sup>-1</sup>) during late the NE MoN period. 374

#### 376 **3.3. Spring Intermonsoon**

#### 377 3.3.1. Wind, Temperature and Salinity

During this season the wind was blowing from the northeast and moving along he 378 east coast of India (Fig. 2c), the wind speed during this period was found to be  $<5ms^{-1}$ . 379 380 This period is the transitional period during which *Trichodesmium* spp. and *Noctiluca* 381 *miliaris* co-existed. These blooms were related to the deepening of the thermocline in this 382 area (Fig. 3c). During this season, the nitrogenous nutrients were low at the surface 383 which varied from 0.0 to 0.39  $\mu$ M (Figs. 4a & 5c). During this time, high (1.66 – 3.38 384  $\mu$ M) silicate content was noticed (Fig. 5c), which suggests an inadequate uptake and 385 coincides with the low density of diatom.

386

#### 387 3.3.2. Chlorophyll a and Trichodesmium distribution

During the spring IntM the Chl *a* highly ranged from  $0.074 - 5.905 \text{ mgm}^{-3}$ . The 388 average Chl a was  $0.759\pm1.025$  mgm<sup>-3</sup> at the surface and  $14.749\pm13.828$  mgm<sup>-2</sup> in the 389 euphotic zone (Fig. 6a). At the surface high Chl *a* was found at  $20^{0}$ N  $70^{0}$ E and low at 390 15<sup>0</sup>28.76N. The total phytoplankton cell counts ranged from 0.025-41.992 cell nos. x 391  $10^{4}L^{-1}$ . The average cell counts were 2.033 $\pm$ 7.302 cell nos. x  $10^{4}L^{-1}$  at the surface and 392  $16.929\pm26.176$  cell nos. x  $10^6$ m<sup>-2</sup> (Fig. 6b) Similarly *Trichodesmium* filaments were also 393 394 high at the same position as the total cell counts. The filaments varied from 0-40.074 nos. x  $10^4$ L<sup>-1</sup> at the surface and 0.021-100.301 nos. x  $10^6$ m<sup>-2</sup>. The average *Trichodesmium* 395 396 filaments were 2.381±8.313 at surface and 6.115±20.737 at the column (Fig. 6c).

In the early springIntM the concentration of Chl *a* varied from 0.169 - 2.741mgm<sup>-3</sup>. The high Chl *a* was found at 19<sup>0</sup>N 66<sup>0</sup>E (2.741 mgm<sup>-3</sup>) was mainly due to the bloom of *Noctiluca miliaris*. The phytoplankton counts ranged from 0.0252 to 1.1164 cell 400 nos. x  $10^{4}L^{-1}$ . The high cell count found at  $19^{0}N$   $66^{0}E$  was also due to the *Noctiluca* 401 *miliaris bloom*. During this month a total of 13 stations were studied out of which only 4 402 stations showed the presence of *Trichodesmium*. The observations indicate that part of 403 the NE Arabian Sea, where the winter cooling affect (low temperature and high 404 nutrients), supports *Noctiluca* and the other part of the Sea, where the nutrient 405 concentration was low and high temperatures prevail, is for *Trichodesmium* area (Table 406 1).

During late spring IntM concentration of Chl *a* varied from 0.074 to 5.905  $\mu$ gL<sup>-1</sup> 407 with the highest values recorded at  $20^{\circ}$ N  $70^{\circ}$ E and the lowest values recorded at  $15^{\circ}$ N 408  $70^{\circ}$ E. Phytoplankton cell counts ranged from 0.033 to 41.991 cell nos. x  $10^{4}$ L<sup>-1</sup>. The 409 phytoplankton cell density showed peak values in coastal waters (21°N 69°E, 20°N 70°E 410 and 15°N 73°E) mostly north of 20°N (Fig. 7c). *Trichodesmium* was the major component 411 of the phytoplankton averaging  $3.8063\pm10.3463$  filaments x  $10^4L^{-1}(n=29)$  with high 412 concentrations (40.0737 filaments x 10<sup>4</sup>L<sup>-1</sup>) at 20<sup>0</sup>N 70<sup>0</sup>E in the coastal waters off 413 Gujarat (Fig. 6c). In general, diatoms and dinoflagellates were less numerous, averaging 414 0.2936 and 0.0630 cells  $L^{-1}$  respectively. However at some locations, like 20<sup>o</sup>N 70<sup>o</sup>E to 415 20<sup>0</sup>N 72<sup>0</sup>E the domination of diatoms was noticed. The dominant diatom, *Nitzschia* sp. 416  $(2.3025 \text{ cells x } 10^4 \text{L}^{-1})$  was found at  $20^0 \text{N} 70^0 \text{E}$ . Dinoflagellates reached a maximum of 417 1.1666 cells x  $10^4 L^{-1}$  at  $20^6 N 70^6 E$ , which was adjacent to the diatom dominated stations 418 indicating succession of groups. The area between 16-18°N in coastal waters was 419 characterized by domination of dinoflagellates. For example, at areas like 16°N 73°E, 420 18°N 72°E and 17°N 72°E dinoflagellates like Ceratium spp., Peridinium spp., and 421 422 Prorocentrum spp. were dominant organisms. At the edge of the shelf, a mixed 423 community of diatoms, dinoflagellates and *Trichodesmium* dominated the flora specially at 21°N 69°E, 20°N 70°E, 15°N 73°E which was in coastal as well as offshore waters 424 (Fig. 6c). The other areas like  $21^{\circ}$ N  $69^{\circ}$ E.  $18^{\circ}$ N  $72^{\circ}$ E and  $17^{\circ}$ N  $72^{\circ}$ E also supported high 425 426 counts of Trichodesmium erythraeum. The Trichodesmium erythraeum was well distributed during this time of year and mostly in succession to the dinoflagellates (south 427 of 18<sup>0</sup>S) and the diatoms (north of 18<sup>0</sup>N). Diatoms were the second most important 428 429 group of organisms that were encountered during the present study, they were also the 430 most diverse in nature.

431

# 432 3.3.3. Primary Production and Nitrogen Fixation

During the spring IntM primary productivity estimated was 1.056-3489.104mgCm<sup>-3</sup>d<sup>-1</sup> which coincided with the *Trichodesmium* bloom area. The average primary productivity was  $255.095\pm668.570$  mgCm<sup>-3</sup>d<sup>-1</sup> (n=32) at the surface and  $2046.857\pm3810.722$  mgCm<sup>-2</sup>d<sup>-1</sup> (n=35) at the euphotic zone (Fig. 8a).

In the early spring IntM, at location  $22^{0}$ N  $66^{0}$ E the *Trichodesmium* was found aggregated at the surface. The primary production due to *Trichodesmium* during this season varied from 0.69 to 0.88 mgCm<sup>-3</sup>d<sup>-1</sup>. The average was found to be 0.79 ± 0.09 mgCm<sup>-3</sup>d<sup>-1</sup> (n=3). About 29 %, out of the total carbon production recorded, was produced by *Trichodesmium*.

In the late spring IntM at  $21^{\circ}$ N  $69^{\circ}$ E and  $15^{\circ}$ N  $73^{\circ}$ E, the high values of Chl *a* and primary productivity were on account of the *Trichodesmium* bloom throughout the water column. *Trichodesmium erythraeum* was the only species found during this season. The average carbon production at surface was  $131.150 \pm 80.67 \text{ mgCm}^{-3}\text{d}^{-1}$  (n=29). Higher 446 rates of primary production in the water column were recorded at locations  $17^{0}N 72^{0}E$ , 447  $21^{0}N 69^{0}E$ ,  $16^{0}N 71^{0}E$  and  $15^{0}N 73^{0}E$ . During this season *Trichodesmium* contributed 448 nearly 30 % to the total primary production.

Nitrogen fixation rate during the spring IntM was reasonably high (271.814±950.263  $\mu$ molNm<sup>-3</sup>d<sup>-1</sup>) at the surface and 720.348±2405.996  $\mu$ molNm<sup>-2</sup>d<sup>-1</sup> in the column (Fig. 8b). The average amount of nitrogen fixed was 1.2±1.1  $\mu$ molNm<sup>-3</sup>d<sup>-1</sup> (n=3). The average nitrogen fixed by *T. thiebautii* was higher (1.75± 0.25) than that fixed by *T. erythraeum*. At 20<sup>0</sup>N 68<sup>0</sup>E and 22<sup>0</sup>N 66<sup>0</sup>E, higher amount of nitrogen was fixed by *T. thiebautii* whereas, at 22<sup>0</sup>N 66<sup>0</sup>E the major contributor to N<sub>2</sub>-fixation was *T. erythraeum*.

456 During late spring IntM the average N<sub>2</sub>-fixed by *T. erythraeum* was 299.81  $\pm$ 457 995.54 µmolNm<sup>-3</sup>d<sup>-1</sup> (n=29). At the surface it varied from 0 to 4580.42 µmolNm<sup>-3</sup>d<sup>-1</sup>. The 458 rate of nitrogen fixation was high at 20<sup>0</sup>N 70<sup>0</sup>E and at column level N<sub>2</sub>-fixated was in the 459 range of about 9.49 to 11471.50 µmol Nm<sup>-3</sup>d<sup>-1</sup>.

460

# 461 **4. Discussion**

The Arabian Sea experiences hydrographic changes, which are related to the monsoon system (Shetye, et al., 1990, Brock and McClain, 1992, Banse and English, 1993, 2000, Madhupratap, et al., 1996). The region of study undertaken fall under the influence of both the NE and the SW monsoon. During the onset of the SW monsoon season, i.e. June, the west coast of India experiences intense upwelling, which is first visible off the southern tip of India. During this period, continental slope waters begin to shoal upward into the shelf region causing substantial cooling, which is observed in the satellite sea surface temperature imagery (Sreejith and Shenoi, 2002; Vinayachandran et
al., 2004). As the season advances the upwelling signatures migrate northward, covering
the entire coast of India by late July (Shetye, et al., 1990). This process is responsible for
bringing in the much needed new nutrients (DeSouza, et al., 1996) that help to trigger
large phytoplankton blooms by early August (Goes, et al., 1992) and continue upto
October.

475 The NE monsoon, marks a period wherein the offshore region receives a 476 substantial amount of new nutrients (Kumar et al., 2009). Satellite derived measurements 477 of Chl a concentrations in the region of the NE Arabian Sea start to increase by 478 December and tend to reach peak values by February (Banse and English, 2000). The 479 data collected during early NE MoN showed that, the nitrate concentrations were undetectable at  $18^{0}05.98$ ,  $18^{0}10.47$ N and  $21^{0}17.81$ N where as low (<0.97 µM) at most of 480 481 the NE Arabian Sea, presumably because a significant fraction of nutrients had already 482 been used by phytoplankton (Parab, et al., 2006), whereas other nutrients such as 483 phosphate and silicate were present. Considering the nutrients enrichment during NE 484 MoN the phytoplankton growth in the water column appears to be reduced by an active 485 grazing by zooplankton (Landry, et al., 1998; Caron and Dennet, 1999) a notion that is 486 reinforced by the elevated level of ammonium concentration found in the water column 487 during early NE MoN estimated at the order of 0.57-6.13 µM of ammonia.

In our study, we found that during low wind speeds ( $< 5 \text{ ms}^{-1}$ ), thick blooms of *Trichodesmium* spp. were observed (Fig. 2). However wind speed >6 knot resulted in disapperance of the *Trichodesmium* bloom. Earlier studies (Chang et al., 2000) have reported that besides wind and water temperature, low nutrients were also factors that 492 favoured growth of the Trichodesmium. In the Arabian Sea, during the non monsoonal 493 period, the upper layers of the water were devoid of nutrients. In such a situation, 494 Trichodesmium had an advantage over other phytoplankton, as it could fix nitrogen and it 495 could make use of phosphate from subsurface layers. The stratification of the water 496 column during low wind conditions (as seen during mid and late NE MoN and early 497 spring IntM) was also an important factor for the Trichodesmium blooms to develop and 498 persist in the Arabian Sea. Associations between abundance of Trichodesmium and 499 vertical stratification have been reported in the North Pacific sub-tropical gyre (Capone, 500 1983). The average mixed layer depth during our study was 10-40m, during 501 *Trichodesmium* bloom.

502 The bloom of Trichodesmium thiebautii observed during the fall intermonsoon 503 continued during the early NE MoN where out of the 15 stations sampled 13 showed the 504 presence of a Trichodesmium thiebautii bloom. The filaments were distributed 505 throughout the water column in the euphotic zone. Deep Chl a maxima was noticed at a 506 Trichodesmium thiebautii concentrations were also observed during mid NE MoN. Out of 507 the 18 stations sampled, 7 stations showed the presence of a mixed population of 508 Trichodesmium erythraeum and Trichodesmium thiebautii. Surface Chl a maxima was observed at 15°14.21N and was largely attributed to Trichodesmium thiebautii and 509 510 Trichodesmium erythraeum.

511 During late NE MoN and early spring IntM, mixed populations of *Trichodesmium* 512 *erythraeum* and *Trichodesmium thiebautii* were noticed at  $20^{0}29.1$ N,  $19^{0}3.7$ N and 513  $15^{0}8.57$ N. However, in the late spring IntM, 29 stations out of the 35 stations sampled 514 supported *Trichodesmium erythraeum* blooms. The rates of primary production during

515 the Trichodesmium bloom were considerably higher than those recorded during non 516 bloom time in the North Atlantic Ocean, which usually ranged from 804 to 1080 mgCm<sup>-</sup>  $^{2}d^{-1}$  (Carpenter, et al., 2004). Several observations have been made in the eastern 517 518 equatorial Atlantic during *Trichodesmium* bloom where Bauerfeind (1987) noted a mean rate of production was 247 mgCm<sup>-2</sup>d<sup>-1</sup> between January and May, 1979 over ten sections 519 between 2°S and 3°N at the 22°W meridian. In March and April of 1989, at 18°N and 520 30<sup>0</sup>W (NW of Cape Verde Islands) primary production over an 11 days period averaged 521 about 500 mgCm<sup>-2</sup>d<sup>-1</sup> (Jochme, et al., 1993). Claustre and Marty (1995) measured 522 primary production at 20<sup>o</sup>N, 31<sup>o</sup>W in autumn 1991 and spring 1992 and where rates 523 averaged  $352 \pm 68$  and  $267 \pm 53$  mgCm<sup>-2</sup>d<sup>-1</sup> respectively. At the oligotrophic EUMELI 524 JGOFS at 22<sup>°</sup>N and 32<sup>°</sup>W, rate of the primary production were ranged from 289 to 276 525 mgCm<sup>-2</sup>d<sup>-1</sup> (Morel, et al., 1996). Steven (1971) reported an annual average of 288 mgCm<sup>-</sup> 526  $^{2}$ d<sup>-1</sup> for a station off the NW coast of Barbados. Beers, et al. (1968) found an annual mean 527 of 110 mgCm<sup>-2</sup>d<sup>-1</sup> for a station in the Caribbean off Jamaica. Chang, et al. (2000) 528 reported that *Trichodesmium* contribution was in the range of 0.003 to 2.27 % of primary 529 530 productivity in the Kuroshio.

We have observed difference in the contributions made by *Trichodesmium erythraeum* and *Trichodesmium thiebautii* towards the organic productivity during our study. In the fall inter-monsoon season about 12 % of total surface primary production was due to only *T. thiebautii* whereas during the spring inter-monsoon it was 30 % because of the presence of *Trichodesmium erythraeum*. Capone, et al., 1998 reported 23 % of total surface primary production was due to *Trichodesmium erythraeum*. The explanation for the high primary production rates could be partially due to the NE monsoon, which generates substantial convective mixing, resulting in deep mixed layers
in the northern Arabian Sea and growth of diatom population (Lee, et al., 2000; Weller,
et al., 1998).

541 Both biomass and productivity of *Trichodesmium* showed large seasonality. Primary productivity due to *Trichodesmium* was as low as 8-15 mgCm<sup>-3</sup>d<sup>-1</sup> during NE 542 MoN ( $8-15 \text{ mgCm}^{-2}\text{d}^{-1}$ ). Another important factor for the success of *Trichodesmium* in 543 544 the Arabian Sea may be the iron input brought about through upwelling and /or 545 atmospheric deposition and land runoff (Witter, et al., 2000). In the present study, large 546 seasonality was also observed in the nitrogen fixation. High N<sub>2</sub>-fixation rate values were recorded during the fall IntM (1108  $\pm$  3511  $\mu$ molNm<sup>-2</sup>d<sup>-1</sup>; n=13) followed by spring IntM 547  $(720 \pm 2406 \text{ }\mu\text{molNm}^{-2}\text{d}^{-1}; \text{ }n=32)$ , however lower values were observed during NE MoN 548  $(24 \pm 32 \text{ } \mu\text{molNm}^{-2}\text{d}^{-1}; \text{ } n=28)$ . Previous studies on the rates of N<sub>2</sub>-fixation by 549 550 Trichodesmium in the sub-tropical North Atlantic were generally much lower (< 10 µmolNm<sup>-2</sup>d<sup>-1</sup>) typically because of the low biomass of the *Trichodesmium* spp. 551 552 (Carpenter, et al., 2004). Published report (Goering, et al., 1966) did indicate the volume 553 specific rates of N<sub>2</sub>-fixation from two cruises off the north-eastern coast of South 554 America. However, they did not provide areal estimates, on the other hand, in southwestern N. Atlantic they reported the mean rate to be  $108 \mu molNm^{-2}d^{-1}$ . Carpenter and 555 556 Romans (1991) postulated rates of N<sub>2</sub>-fixation for Trichodesmium of 710 to 3600  $\mu$ molNm<sup>-2</sup>d<sup>-1</sup> based on the high abundance of *Trichodesmium* encountered in the tropical 557 558 portion of a transect through the Atlantic. In contrast to the above studies, Lipschultz and 559 Owens (1996) produced a much smaller estimate for the contribution of Trichodesmium in the same region. Capone, et al. (2005) reported the average N<sub>2</sub>-fixation rate (898 560

 $\mu$ molNm<sup>-2</sup>d<sup>-1</sup>) in the same study area. In the Southern East China Sea, Chang, et al. 561 562 (2000) reported that the N<sub>2</sub>-fixation rate varied from 0.1 to 59.4  $\mu$ molNm<sup>-2</sup>d<sup>-1</sup>. This was 563 much lower than the results obtained during the our study in the NE Arabian Sea (Table 564 2). The Arabian Sea is known as an active site for nitrogen fixation (Devassy, et al., 1978, Capone, et al., 1998). One factor that favors the establishment of N<sub>2</sub>-fixers in the 565 566 Arabian Sea is the  $NO_3$  deficit arising from local and far-field denitrification. The  $NO_3$ :  $PO_4^{3-}$  ratio was lower than 15 at all depths in the Arabian Sea (Morrison, et al., 1998). As 567 568 the N- deficient sub-surface waters were brought to the surface through upwelling and 569 vertical mixing, the nutrients were taken up by the phytoplankton in a (Redfield) ration of 15 (by atoms), and this led to a surplus of  $PO_4^{3-}$  in surface waters. This excess  $PO_4^{3-}$  is 570 observed most strikingly within the coastal anoxic zone where  $PO_4^{3-}$  concentrations in 571 excess of 3  $\mu$ M occur when denitrification removes all dissolved NO<sub>3</sub><sup>-</sup>. The organisms 572 that fix N<sub>2</sub> also require relatively large amount of iron (Fe), which could be supplied by 573 574 atmospheric deposition, mobilization from shelf and slope sediments and upwelling. 575 Consequently, blooms of *Trichodesmium* spp. occur over large parts of the Arabian Sea 576 mostly during the spring inter-monsoon season when surface waters are  $NO_3^-$  depleted. 577 The calmness of the sea during this season allows the establishment of warm surface 578 waters, reduced turbulance and high euphotic zone light intensities, conditions that also 579 favour nitrogen fixation (Devassy, et al., 1978, Capone, et al., 1997, Capone, et al., 580 1998). Apart from the experimentally measured N<sub>2</sub>-uptake rates (Dugdale, et al., 1964, 581 Capone, et al., 1998), evidence for  $N_2$ -fxation in the Arabian Sea has been provided by 582 data on natural isotope abundance in NO<sub>3</sub><sup>-</sup> (Brandes, et al., 1998). The estimated N<sub>2</sub>- fixation rate for the region is 3.3 TgNyr<sup>-1</sup> (Bange, et al., 2000) which is speculated to play
an important role in understanding biogeochemistry of the region.

Thus our present study illustrates the seasonal distribution of *Trichodesmium* and its ability to fix carbon and nitrogen in the Arabian Sea. A large part of the Arabian Sea tends to become oligotrophic during a major part of the year (6-7 months) where primary productivity is impacted by the nitrogen fixed by the *Trichodesmium*.

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**Insert Table 2 here** 

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593 During the present study, a total of 143 stations were sampled in the NE Arabian 594 Sea in order to study the *Trichodesmium* bloom. Majority of these stations (93 stations) 595 showed the presence of Trichodesmium i.e. 65 % of the total stations were represented 596 by Trichodesmium. A significant finding of the present study was the occurrence of T. 597 thiebautii in the offshore part of the Arabian Sea, just after the SW monsoon, as early as November (Fig. 9). This was largely attributed to the low temperatures (22-27<sup>0</sup>C) and 598 599 low nitrate levels in the euphotic zone; Trichodesmium erythraeum began appearing in 600 February in the coastal water and spread to offshore water during the months of April and 601 May, when temperatures generally high and nitrate concentrations are low and so 602 Trichodesmium ervthraeum thrives in this environment.

603 Insert Fig. 9 here

Another important feature observed during the NE MoN was that *Trichodesmium thiebautii* occurred as deep as 70 – 80 m in the water column. Phosphate availability may be an important factor for *Trichodesmium* to appear (Karl, et al., 1992). In the NE Arabian Sea, the contribution of *Trichodesmium* to total phytoplankton counts varied

from 1% to 95 %, with the highest counts being observed in the spring IntM. The dominant phytoplankton species in the *Trichodesmium* bloom were *Trichodesmium erythraeum*, *Trichodesmium thiebautii*, *Navicula* spp., *Nitzschia closterium*, *Nitzschia seriata*, *Rhizosolenia setigera*, *Rhizosolenia shrubsolei*, *Chaetoceros curvisetus*, *Noctiluca miliaris*, *Melosira moniliformis*, *Chaetoceros decipiens*, *Rhizosolenia hebetata* and *Thalassiothrix frauenfeldii*.

614 Another noteworthy observation made during the present investigation was the 615 occurrence of mixed blooms of Trichodesmium erythraeum and Noctiluca miliaris from 616 late NE MoN and early spring IntM in the shelf region, whereas, in the offshore region, 617 Noctiluca dominated the bloom, this coincided with a reduction in the density of the 618 Trichodesmium filaments. Further, unique occurrence of isolated patches of decomposed 619 filaments in combination with *Noctiluca* at a few stations coupled with elevated nutrient 620 suggests that, the post Trichodesmium bloom triggers the Noctiluca miliaris bloom in the 621 Arabian Sea. It is a well known fact that the Arabian Sea becomes oligotrophic during 622 the late NE MoN, which supports the occurrence of a *Noctiluca miliaris* (Gomes et al., 623 2008; 2009). Winter convection (February-March) that brings high nutrients to surface 624 water in the NE Arabian Sea triggers blooms of the *Noctiluca miliaris*. Subsequently, in 625 the Arabian Sea Noctiluca and Trichodesmium blooms coexist with high counts of 626 Navicula sp., Nitzschia closterium, Nitzschia seriata and Thalassiosira subtilis. However 627 Trichodesmium spp. which contributes 0.11 to 60 % of total primary production of the 628 Arabian Sea remains important.

It appears that a sizable portion of nitrogen will be brought back to the surface
layer, (may be N0<sub>3</sub>), which could support phytoplankton productivity (new production)

631 in the Arabian Sea. On analyzing the data we collected from all the stations, we found 632 that the organic carbon produced by Trichodesmium in the region was about 0.263TgCvear<sup>-1</sup>, this value was, in fact very close to the amount of nitrogen fixed by the 633 634 Trichodesmium in the region during same period, which was estimated to be about 0.2976 TgNyear<sup>-1</sup>. At times large amounts of biomass get accumulated at surface during 635 Trichodesmium bloom in the Arabian Sea. This cyanobacteria is effective for N<sub>2</sub> fixation, 636 637 but not as primary producer and hence is responsible for the matching carbon fixation and 638 nitrogen fixation quantity ratio nearing one. Thus the presence of Trichodesmium spp. 639 recorded from November to April in the Arabian Sea (Fig. 9) is a important factor 640 responsible for maintaining the ecosystem of the Arabian Sea during oligotrophic 641 conditions.

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645

# 644 **5. Conclusions**

646 The two species of the Trichodesmium form blooms in the Arabian Sea, these 647 blooms depend upon environmental conditions. Accordingly blooms appear in the coastal 648 water or open ocean. Trichodesmium thiebautii blooms were noticed just after monsoon 649 in the low nutrient offshore environment and low temperature. Similar conditions extend 650 in the coastal water during January where T. thiebautii was noticed. March onwards 651 offshore waters are marked with T. erythraeum blooms and April onwards coastal waters. These blooms synthesize 0.263TgCyear<sup>-1</sup> and fix 0.2976 TgNyear<sup>-1</sup>. The nitrogen fixed 652 653 help the productivity of the Arabian Sea which is otherwise limited due to the depletion 654 of nitrogenous nutrients, between the two monsoon systems.

## 656 Acknowledgment

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# 663 **References**

- 664 665
- Bange, H. W., Naqvi, S. W., Codispoti, L. A., 2005. The nitrogen cycle in the
  Arabian Sea. Progress of Oceanography 65, 145 158.
- Bange, H. W., Rixen, T., Johansen, A. M., Siefert, R. L., Ramesh, R., Ittekkot, V.,
  Hoffmann, M R., 2000. A revised nitrogen budget for the Arabian Sea. Global
  Biogeochemical Cycles 14, 1283 1297.
- Banse, K., English, D. C., 2000. Geographical differences in seasonality of CZCSderived phytoplankton pigment in the Arabian Sea for 1978-1986. Deep-Sea
  Research II 47, 1623-1677.
- 4. Banse, K. English, D. C., 1993. Revision of satellite-based phytoplankton pigment data from the Arabian Sea during the northeast monsoon. Marine Research (Pakistan) 2, 83 101.
- 5. Barber, R. T., Marra, J., Bidigare, R. R., Codispoti, L. A., Halpern, D., Johnson, Z.,
  Latasa, M., Goericke, R., Smith, S. L., 2001. Primary productivity and its regulation
  in the Arabian Sea during 1995. Deep-Sea Research II 48, 1127-1172.
- 680
   6. Basu, S., Matondkar, S. G. P., Furtado, I., 2011. Enumeration of bacteria from a *Trichodesmium* spp. bloom of the eastern Arabian Sea: Elucidation of their possible role in biogeochemistry. Journal of Applied Phycology 23, 309-319.
- 683 7. Bauerfeind, E., 1987. Primary production and phytoplankton biomass in the equatorial region of the Atlantic at  $22^{0}$ W.Oceanologia acta 10, 131 136.
- 8. Beers, J. R., Steven, D. M., Lewis, J. B., 1968. Primary productivity in the
  Caribbean Sea off Jamaica and the tropical North Atlantic off Barbados. Bulletin of
  Marine Science 18, 86 104.
- Bhattathiri, P. M. A., Pant, A., Sawant, S. S. Gauns, M., Matondkar, S. G. P., 1996.
  Phytoplankton production and chlorophyll distribution in the eastern and central
  Arabian Sea in 1994-1995. Current Science 71, 857-862.
- 691 10. Bidigare, R. G., Charles, C. T., 2002. HPLC phytoplankton pigments: sampling,
   692 laboratory methods and quality assurance procedures. In: Fargion, G. S., Muellier,

- J. L (Eds.), Protocols For Satellite Ocean Colour Validation Revision. NASA,
  Goddard Space Flight Centre, Greenbelt, Maryland 20771, pp. 154-160.
- 11. Brandes, J. A., Devol, A. H., Jayakumar, D. A., Yoshinary, T., Naqvi, S. W. A.,
  1998. Isotopic composition of nitrate in the Central Arabian Sea and eastern tropical
  North Pacific: a tracer for mixing and nitrogen cycle. Limnology and Oceanography
  43, 1680 1689.
- Brock, J. C., McClain, C. R., 1992. Interannual variability of the southwest
  monsoon phytoplankton bloom in the northwest Arabian Sea. Journal of
  Geophysical Research 97, 733-750.
- 13. Capone, D. G., Subramaniam, A., Montoya, J. P., Voss, M., Humborg, C.,
  Johansen, A. M., Siefert, R. L., Carpenter, E. J.1998. An extensive bloom of the N<sub>2</sub>fixing cyanobacterium, *Trichodesmium erythraeum* in the central Arabian Sea.
  Marine Ecology Progress Series 172, 281-292.
- 14. Capone, D. G., Burns, J. A., Montoya, J. P., Subramaniam, A., Mahaffrey, C.,
  Gunderson, T., Michaels, A. F., Carpenter, E. J., 2005. Nitrogen fixation by *Trichodesmium* spp: An importanct sources of new nitrogen to the tropical and
  subtropical North Atlantic Oceans. Global Biogeochemical Cycle 19, 1 17.
- Tichodesmium a Globally Significant Cyanobacterium. Science 276, 1221-1229.
- 712 16. Caron, D. A., Dennett, M. R., 1999. Phytoplankton growth and mortality during the
  713 1995 northeast monsoon and spring intermonsoon in the Arabian Sea. Deep-Sea
  714 Research 46, 1665-1690.
- 715 17. Carpenter, E. J., 1983. Nitrogen fixation by marine Oscillatoria (Trichodesmium) in
  716 the world's oceans. In: Carpenter, E. J., Capone, D. G. (Eds.), Nitrogen in the
  717 Marine Environment. Academic Press, New York, pp. 65 103.
- 18. Carpenter, E. J., Price IV, C. C., 1977. Nitrogen fixation, distribution, and
  production of *Oscillatoria (Trichodesmium)* in the northwestern Atlantic Ocean and
  Caribbean Sea. Limnology Oceanography 22, 60-72.
- 19. Carpenter, E. J., Romans. K., 1991. Major role of the Cyanobacterium
   *Trichodesmium* in nutrient cycling in the North Atlantic Ocean. Science 254, 1356 1358.
- 20. Carpenter, E. J., Subramaniam, A., Capone, D. G., 2004. Biomass and primary
  productivity of the cyanobacterium *Trichodesmium* spp. in the tropical N Atlantic
  Ocean. Deep-Sea Research 51, 173-203.
- 21. Chang, J., Chiang, K. P., Gong, G. C., 2000 Seasonal variation and cross-shelf
  distribution of the nitrogen-fixing cyanobacterium, *Trichodesmium*, in southern
  East China Sea. Continental Shelf Research 20, 479-492.
- 22. Chen, Y. L. L., Chen, Y., Lin, Y. H., 2003. Distribution and downward flux of
   *Trichodesmium* in the South China Sea as influenced by the transport from the
   Kuroshio current. Marine Ecology Progress Series 259, 47 57.

- 23. Claustre, H., Marty, J. C., 1995. Specific phytoplankton biomasses and their
  relation to primary production in the tropical north Atlantic. Deep-Sea Research 42,
  1475-1493.
- 24. Desa, E. S., Suresh, T., Matondkar, S. G. P., Mascarenhas, A., Desa, E., Goes, J.,
  Parab, S. G. P., Shaikh, N., Fernandes, C. F. C., 2005. Detection of *Trichodesmium*bloom patches by IRS-P4/OCM ocean colour sensor and by in situ measurements
  along the Eastern Arabian Sea Coast. Indian Journal of Marine Science 34, 374386.
- 741 25. DeSouza, S. N., Kumar, M. D., Sardessai, S. V., Sarma, V. S., Shirodkar, P. V.,
  742 1996. Seasonal variability in oxygen and nutrients in the central and eastern
  743 Arabian Sea. Current Science 71, 857-862.
- 26. Devassy, V. P., Bhattathiri, P. M. A. and Qasim, S. Z. (1978) *Trichodesmium*phenomenon. Indian Journal Marine Science 7, 168-186.
- 746 27. Dugdale, R. C., Goering, J. J., 1967. Uptake of new and regenerated forms of
  747 nitrogen in primary productivity. Limnology and Oceanography 12, 677 680.
- 28. Dugdale, R. C., Goering, J. J., Ryther, J. H., 1964. High nitrogen fixation rates in the Sargasso Sea and the Arabian Sea. Limnology and Oceanography 9, 507 – 510.
- 29. Eleuterius, L., Perry, H., Eleuterius, C., Warren, J., Caldwell, J., 1981. Causative
  analysis on a nearshore bloom of Oscillatoria erythraea (*Trichodesmium*) in the
  northern Gulf of Mexico. Northeast Gulf Science 5, 1-11.
- 30. Eppley, R.W., Peterson, B. J., 1979. Particulate organic matter flux and planktonic
  new production in the deep ocean. Nature 282, 677–680.
- 31. Fernandez, A., Mourino-Carballido, B., Bode, A., Varela, M., Maranon E., 2010.
  Latitudinal distribution of Trichodesmium spp. and N2 fixation in the Atlantic
  Ocean. Biogeosciences Discuss. 7, 2195-2225.
- 32. Goering, J. J., Dugdale, R. C., Menzel, D. W., 1966. Estimates of in situ rates of nitrogen uptake by Trichodesmium sp. in the tropical Atlantic ocean. Limnology and Oceanography 11, 614-620.
- 33. Goes, J. I., Gomes, H. do R., Kumar, A., Gouveia, A., Devassy, V.P., Parulekar, A.
  H., Rao, L.V.G., 1992. Satellite and ship studies of phytoplankton along the west
  coast of India. In: Desai, B.N. (Eds.,), Oceanography of the Indian Ocean. Oxford
  and IBH Publishing Co., New Delhi, pp. 67–80.
- 34. Gomes, H.do.R., Goes, J. I., Matondkar S. G. P., Parab, S. G., Al-Azri, A. R. N.,
  Thoppil, P. G., 2008. Blooms of *Noctiluca miliaris* in the Arabian Sea-An *in situ*and satellite study. Deep- Sea Research I 55, 751-765.
- 35. Gomes, H. do R., J. I. Goes, S. G. P. Matondkar, Parab, S. G., A. Al-Azri, Thoppil,
  P. G., 2009. Unusual Blooms of the Green *Noctiluca Miliaris* (Dinophyceae) in the
  Arabian Sea during the Winter Monsoon. In: Wiggert, J. D., Hood, R. R., Naqvi, S.
  W. A., Brink K. H., Smith S. L. (Eds.,), Indian Ocean Biogeochemical Processes
  and Ecological Variability. AGU Geophysical Monograph Series 185. American
  Geophysical Union, Washington, DC, pp. 347-363.

- 36. Gunderson, K. R., 1976. Structure and biological dynamics of the oligotrophic
  ocean photic zone of the Hawaiian Islands. Pacific Science 30, 45-68.
- 37. Hegde, S., Anil, A. C., Patil, J. S., Mitbavkar, S., Krishnamurthy, V., Gopalkrishna,
  V. V., 2008. Influence of environmental settings on the prevalence of
  Trichodesmium spp. in the Bay of Bengal. Marine Ecology Progress Series 356,
  93-101.
- 38. Holl, C. M., Villareal, T. A., Payne, C. D., Clayton, T. D., Hart, C., Montoya, J.,
  2007. Trichodesmium in the western Gulf of Mexico: 12N2-ixation and natural
  abundance stable isotope evidence. Limnology and Oceanography 52, 2249-2259.
- 39. Jochem, F. J., Pollehne, F., Zeitzschel, B., 1993. Productivity regime and
  phytoplankton size structure in the Arabian Sea. Deep-Sea Research 40, 711-735.
- 40. Jyothibabu, R., Madhu, N. V., Murukesh, N., Haridas, P. C., Nair, K. K. C.,
  Venugopal, P., 2003. Intense blooms of *Trichodesmium erythraeum* (Cyanophyta)
  in the open waters along east coast of India. Indian Journal of Marine Science 32,
  165-167.
- 41. Karl, D., Letelier, R., Hebel, D., Bird, D., Winn, C., 1992. *Trichodesmium* blooms and new nitrogen in the North Pacific Gyre. In: Carpenter, E. J., Capone, D. G., Rueter, J. G. (Eds.,), Marine Pelagic Cyanobacteria: *Trichodesmium* and other Diazotrophs. Kluwer Academic, pp. 219-237.
- 42. Karl, D., Letelier, R., Tupas, L., Dore, J., Christian, J., Hebel, D., 1997. The role of
  nitrogen fixation in biogeochemical cycling in the subtropical North Pacific Ocean.
  Nature 388, 533-538.
- Kitajima, S., Furuya, K., Hashihama, F., Takeda, S., 2009. Latitudinal distribution
  of diazotrophs and their nitrogen fixation in the tropical and subtropical western
  North Pacific. Limnology and Oceanography 54, 537-547.
- 44. Kumar, S. P., Narvekar, J., Nuncio, M., Gauns, M., Sardesai, S., 2009. What drives
  the biological productivity of the Northern Indian Ocean? In: In: Wiggert, J. D.,
  Hood, R. R., Naqvi, S. W. A., Brink K. H., Smith S. L. (Eds.,), Indian Ocean
  Biogeochemical Processes and Ecological Variability. AGU Geophysical
  Monograph Series 185. American Geophysical Union, Washington, DC, pp.33-56.
- 45. Landry, M.R., Brown, S.L., Campbell, L., Constaninou, J. and Liu, H. (1998)
  Estimating rates of growth and microzooplankton grazing in the Arabian Sea during
  monsoon forcing. Deep- Sea Research II 45, 2353–2368.
- 46. Lee, C.M., Jones, B.H., Brink, K.H., Fischer, A.S., 2000. The upper-ocean response
  to monsoonal forcing in the Arabian Sea: seasonal and spatial variability. Deep-Sea
  Research 47, 1177–1226.
- 47. Letelier, R. M., Karl, D. M., 1996. Role of *Trichodesmium* spp. in the productivity
  of the subtropical North Pacific Ocean. Marine Ecology Progress Series 133, 263273.
- 48. Lewis, M. R., Harrison, W. G., Oakey, N. S., Hebert, D., Platt, T., 1986. Vertical
  nitrate fluxes in the oligotrophic ocean. Science 234, 870 –873.

817 49. Lipschultz, F., Owens, N. J. P., 1996. An assessment of nitrogen fixation as a 818 source of nitrogen to the North Atlantic Ocean. Biogeochemistry 35, 261 - 274. 819 50. Lugomela, C., Lyimo, T. J., Bryceson, I., Semesi, A. K., Bergman, B., 2002. 820 Trichodesmium in coastal waters of Tanzania: diversity, seasonality, nitrogen and 821 carbon fixation. Hydrobiologia 477, 1-13. 822 51. Madhupratap, M., Prasanna Kumar, S., Bhattathiri, P.M.A., Dileep Kumar, M., 823 Raghukumar, S., Nair, K.K.C., Ramaiah, N., 1996. Mechanism of the biological 824 response to winter cooling in the northeastern Arabian Sea. Nature 384, 549–552. 825 52. Morel, A., Antoine, D., Babin, M., Dandonneau, Y., 1996. Measured and modeled 826 primary production in the Northeast Atlantic (EUMELI JGOFS program): The 827 impact of natural variations in photosynthetic parameters on model predictive skill. 828 Deep-Sea Research I 43, 1273-1304. 829 53. Morrison, J. M., Codispoti, L. A., Gaurin, S., Jones, B., Manghnani, V., Zheng, Z., 830 (1998). Seasonal variations of hydro-graphic and nutrient filed during the US 831 JGOFS Arabian Sea process study, Deep-Sea Research II 45, 2053-2101. 832 54. Mulholland, M. R., Bernhardt, P. W., Heil, C. A., Bronk, D. A., O'Neil, J. H., 2006. 833 Nitrogen fixation and release of fixed nitrogen by Trichodesmium spp. in the Gulf 834 of Mexico. Limnology and Oceanography 51, 1762-1776. 835 55. Mulholland, M. R., Capone, D. G., 2009. Dinitrogen Fixation in the Indian Ocean. 836 In: Wiggert, J. D., Hood, R. R., Naqvi, S. W. A., Brink K. H., Smith S. L. (Eds.,), 837 Indian Ocean Biogeochemical Processes and Ecological Variability. AGU 838 Geophysical Monograph Series 185. American Geophysical Union, Washington, DC, pp. 167-186. 839 840 56. Naqvi, S. W., Moffett, J. W., Gauns, M. U., Narvekar, P. V., Pratihary, A. K., Naik, 841 A., Shenoy, D. M., Jayakumar, D. A., Goepfert, T. J., Patra, P. K., Al-Azri, A., 842 Ahmed, S. I., 2010. The Arabian Sea as a high-nutrient, low-chlorophyll region 843 during the late Southwest Monsoon. Biogeosciences 7, 2091-2100. 844 57. Orcutt, K. M., Lipschultz, F., Gundersen, K., Arimoto, R., Michaels, A. F., Knap, 845 A.H., Gallon, J. R., 2001. A seasonal study of the significance of N<sub>2</sub> fixation by 846 Trichodesmium spp. at the Bermuda Atlantic Time-series Study (BATS) site. 847 Deep-Sear Research II 48, 1583-1608. 848 58. Parab, S. G., Prabhu Matondkar, S. G., Gomes, H.do. R., Goes, J. I., 2006. 849 Monsoon driven changes in phytoplankton population in the Eastern Arabian Sea as 850 revealed by microscopy and HPLC pigment analysis. Continental Shelf Research 851 26, 2538-2558 852 59. Platt, T., Jauhari, P., Satyendranath, S., 1992. The importance and measurement of 853 new production, in primary productivity and biogeochemical cycles in the sea. In: 854 Falkowski, P. G., Woodhead, A. D., (Eds.), Springer, New York, pp. 273 – 384. 60. Qasim, S. Z., 1970. Some characteristics of a Trichodesmium bloom in the 855 Laccadives. Deep-Sea Research II 17, 655-660. 856

- 857 61. Ramamurthy, V. D., Bhargava, R. A., Selvakumar, R. A., 1972. Studies on the
  858 blooms of *Trichodesmium erythraeum* (EMR) in the waters of the central west coast
  859 of India. Current Science 41, 803-804.
- 860 62. Saino, T., 1977. Biological Nitrogen Fixation in the Ocean with Emphasis on the
  861 Nitrogen Fixing Blue-Green Alga, *Trichodesmium* and its significance in the
  862 Nitrogen Cycle in the Low Latitude Sea Areas, University of Tokyo. 153.
- 863 63. Sawant, S., Madhupratap, M., 1996. Seasonality and composition of phytoplankton
  864 in the Arabian Sea. Current Science 71, 869-873.
- 865 64. Shetye, S. R., Gouveia, A. D., Shenoi, S. S. C., Soundar D., Michael, G. S.,
  866 Almeida, A. M., Santanam, K., 1990. Hydrography and circulation off the west
  867 coast of India during the Southwest Monsoon 1987. Journal of Marine Research 48,
  868 359-378.
- 869 65. Smith, S. L., 2001. Understanding the Arabian Sea: Reflections on the 1994-1996
  870 Arabian Sea Expedition. Deep-Sea Research II 48, 1385-1402.
- 66. Smith, S. L., 2005. The Arabian Sea of the 1990s; New biogeochemical
  understanding. Progress of Oceanography 65, 113-115.
- 873 67. Sreejith, O. P., Shenoi, S. C., 2002. An evaluation of satellite and in situ based sea
  874 surface temperature data sets in the North Indian Ocean region. International
  875 Journal of Remote Sensing 23, 2165-2175.
- 876 68. Steven, D. M., 1971. Primary productivity of the tropical western Atlantic ocean
  877 near Barbados. Marine Biology 10, 261 264.
- 878 69. Subramaniam, A., Carpenter, E. J., 1994. An empirically derived protocol for the
  879 detection of blooms of the marine cyanobacterium *Trichodesmium* using CZCS
  880 imagery. International Journal of Remote Sensing 15, 1559-1569.
- 70. Tomas, C. R., 1997. Identifying Marine Phytoplankton. Academic Press, New
  York, USA, pp. 858.
- 71. UNESCO, (1994). Protocols for the Joint Global Ocean Flux Study (JGOFS) core
  measurements, Manual and guides 29, Scientific Committee on Oceanic Research,
  pp. 170.
- Vinayachandran, P. N., Chauhan, P., Moohan, M., Nayak S., 2004. Biological
  responses of the sea around Sri Lanka to summer monsoon. Geophysical Research
  Letters 31 LO1302, doi:10.1029/2003GL018533.
- 73. Weller, R. A., Baumgartner, M. F., Josey, S. A., Fischer, A. S., Kindle, J. C. 1998.
  Atmospheric forcing in the Arabian Sea during 1994-1995: observations and comparisons with climatology and models. Deep-Sea Research II 45, 1961 1999.
- 892 74. Witter, A. E., Lewis, B. L., Luther, G. W., 2000. Iron speciation in the Arabian Sea.
  893 Deep-Sea Research II 47, 1517-1539.
- 894 75. Wright, S.W., Jeffrey, S. W., Mantoura, R. G. C., Llewellyn, C. A., Bjornland, D.
  895 R., Welschmeyer, N., 1991. Improved HPLC method for the analysis of

chlorophylls and carotenoids from marine phytoplankton. Marine Ecology Progress
Series 77, 183-196.

898 899 900 901	List of Tables					
902	1.	Dominant phytoplankton in the Arabian Sea				
	2.	Historical data of N <sub>2</sub> fixation by <i>Trichodesmium</i> studies in the world ocean				
903						

# Table(s)

		Surface		Column			
Fall	Trichodesmium	Nitzschia	Amphora	Chaetoceros	Melosira	Chaetoceros	
IntM	thiebautii	Closterium	Hyaline	socialis	moniliformis	Danicus	
NE MoN	Gymnodinium gracile	Prorocentrum minimus	Noctiluca milaris	Trichodesmium thiebautii	Nitzschia longissima	Noctiluca miliaris	
Spring	Trichodesmium	Noctiluca	Coscinodiscus	Navicula	Trichodesmium	Rhizosolenic	
IntM	erythraeum	milaris	marginatus	delicatula	erythraeum	stolterfothii	

# **Table 1** Dominant phytoplankton in the Arabian Sea.905

Location	Sampling period	Mean Rate (µmolN m <sup>-2</sup> d <sup>-1</sup> )	SE	n	Method	Source
Southwestern N	Nov 1964	41	7.6	19	$^{15}N_{2}$	Goering et al 1966
Atlantic $0-24^{\circ}N$ , 45-66 $^{\circ}N$	May 1965	108	24	17	uptake	Goornig et al., 1900
Carribbean, 12-22 <sup>0</sup> N		161	20	12	AR	Carpenter and Price, 1977
Bermuda Atlantic Time-Series studies	1995-1997	41		14	<sup>15</sup> N <sub>2</sub> uptake	Orcutt et al., 2001
N. Pacific, 21 <sup>0</sup> N, 159 <sup>0</sup> W	1972	134		2	ĀR	Gunderson, 1976
East China Sea, 10- 25 <sup>0</sup> N		126	49	32	AR	Saino, 1977
Hawaii Ocean Time- Series/station ALOHA	1990-1992	84	43	8	AR	Karl et al., 1997
Arabian Sea, 7-10 <sup>0</sup> N	May 1995	35	7.4	9	AR	Capone et al., 1998
Arabian Sea, 10 <sup>0</sup> N bloom	May 1995	99	25	9	AR	Capone et al., 1998
Coastal Tanzania (bloom-upper 0.5m)	1975-1999	59			AR	Lugomela et al., 2002
Gulf of Mexico	July 2000	85	18	4	<sup>15</sup> N <sub>2</sub> uptake	Holl et al., 2007
Western North Pacific 0-12 <sup>0</sup> N 16-24.5 <sup>0</sup> N	Nov-Dec 2007 May-June 2008	538 939		4 3	ĀR	Kitajima at al., 2009
Atlantic Ocean $30^{0}$ N to $30^{0}$ S	Nov-Dec 2007 April-May 2008				AR	Fernandez et al., 2010
North Gyre,		25				
Equator		66				
South Gyre		3				
Arabian Sea $(12-22^{\circ}N \text{ and } 65-72^{\circ}E)$						
Fall Intermonsoon	Nov 2001	1108.15	973.8	13	Trichome	Present study
NE monsoon	Jan-Feb 2003,2004; Dec 2004	24.75	5.9	28	based (Capone	Present study
Spring	April 2000	720.35	425 3	32	1998)	Present study
Intermonsoon	March 2003, 2004	120.30	120.0	52	1770)	1 robont study

# 910 **Figure captions**

911

1	Study area	chowing	compling	nositions
1.	Study area	showing	sampning	positions

- The pattern of wind direction during (a) fall intermonsoon (b) NE monsoon and
   (c) Spring intermonsoon.
- Variations in water temperature (<sup>0</sup>C) and salinity (psu) during (a) fall intermonsoon (b) NE monsoon and (c) Spring intermonsoon.
- 4. Variations in nutrients during different seasons (a) Nitrate (b) Nitrite (c)Phosphate and (d) Silicate.
- 5. Vertical profiles of nutrients during (a) fall intermonsoon (b) NE monsoon and(c) Spring intermonsoon.
- Seasonal distribution of (a) Chlorophyll *a* (b) Phytoplankton cell density and (c)
   *Trichodesmium* filaments. The sets of panels at the left side are surface and right side are column.
- 7. Seasonal distribution of (a) Primary Productivity and (b) Nitrogen fixation during *Trichodesmium* bloom. The sets of panels at the left side are surface and right side are column.
- 8. Depth wise distribution of Chlorophyll *a* and Phytoplankton cell counts (including *Trichodesmium* spp.) during *Trichodesmium* bloom.
- 9. Chlorophyll *a* from OCM in the Arabian Sea during different months.





926 Fig. 2







Fig. 4 









1001 Fig. 8



- Fig. 9