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HEAT BUDGET OF A SHORE POLYNYA, DRONNING MAUD LAND COAST, ANTARCTICA DURING AUSTRAL SUMMER SEASON, 1987-88

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1. INTRODUCTION

an isolated water body Polynya, an isolated wate surrounded by ice, influences local considerably. climate The oceaninteraction over the polynya atmosphere the heat losses helps in increasing several times higher than from the region which is covered with ice. Amongst the Antarctic polynyas, the Weddel polynya mostly drawn the attention ο£ has investigators and some suggestions have been proposed for its evolution proposed 1982). (Carsey, 1980; Gordon, East οf Weddel Sea, smaller shore-type polynyas are also known to exist especilly along the Dronning Maud Land coast. Ever since the first Indian Antarctica expedition of 1982, a coastal polynya off the Indian ice shelf station 'Dakshin Gangotri' become a regular visiting place for supply vessels which bring exped the vessels which bring expedition members/material to the icy continent. the studies on biological Some productivity and chemical aspects of this polynya have been made by Sengupta and Qasim (1982) and Pant (1986). Detailed investigations on heat exchange between the ocean and atmosphere have been carried out more for Arctic polynyas (Smith, 1981; Topham et al, 1983 and Smith et al 1983) as compared to those near Antartica. tidal flows across shallow regions in the Arctic archipelago are responsible for stronger mixing causing the surface waters to be warm and remain above the freezing point. These warm waters thereby enhance sensible heat flux which is found to be more than 50% of the total heat loss (Harteg et al, 1983).

Recent investigation of Kumar and Sadhuram (1989) on heat balance of shore polynya in the Queen Maud Land coast of Antarctica has sought the balance of seasonal surface heat gain against the heat loss through oceanic interior processes affecting steady conditions. However, the examination of sea surface temperature (SST) variation in the same polynya during 1987-88 austral summer season has aroused some interest in reinvestigating its heat budget as SST has showed a decreasing trend. We present here the evolution of the heat budget of the polynya with additional information on advection rates using moored current meter data.

2. DATA AND METHODS

Fig.1 shows the area of the polynya off the Dronning Maud Land coast of Anarctica where the ship 'MV THULELAND'

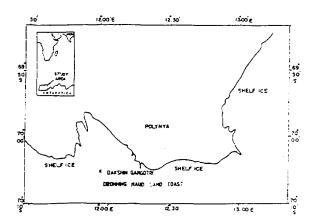


Fig.1 Location of the shore polynya off pronning Maud land Coast, Antarctica.

was remained during the period from 30 December to 28 February, 1988. Necessary information for the computation of various surface heat balance components using bulk-aerodynamic method was taken from the ship's log of surface meteorological observations which were recorded at a 6-hourly interval. Though errors are possible due to non-uniform level of observations and structure of the ship, they have been usually considered not more than 20% (Hastenrath and Lamb, 1979).

3. RESULTS

Daily averages of various heat fluxes are estimated and their evolutions are presented in figures 2 to 5. The variation of incoming global radiation (Q₁) shows, in general, a seasonal decreasing trend during January-February (Fig. 2). High values up to 300 W m occur in the middle of January while the incidence of values as low as 50 W m is seen at the end of February. The effect of cloudiness(N) on Q₁ is quite evident from the variation of N which is seen as the mirror image of Q₁.

Latent heat flux $(Q_{\rm e})$ shows (Fig. 3) a rather steady regime in January whereas in February an increasing trend is seen. The role of vapour pressure difference (e -e) on Q is felt more in January. On the other hand, winds strengthened by about six times, influence $Q_{\rm e}$ in February.

A conspicous result of the present study is the occurrence of sensible heat flux (Q) around the same order as that of evaporative heat flux (Q). This shows Q as an important one in surface heat loss

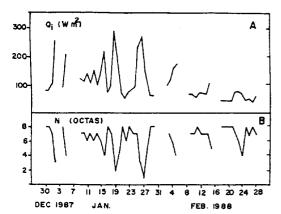


Fig. 2 paily variation of (A) global radiation and (B) cloud amount.

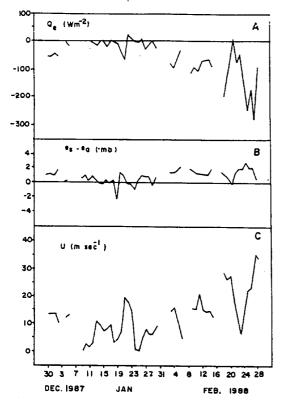


Fig.3 Daily variation of (A) latent heat
flux, (B) vapour pressure difference
and (C) wind speed.

processes (Fig. 4) unlike in tropics where Q_s is not significant. Q_s is mainly positive during January gaining heat due to negative values of sea-air temperature difference (T_s-T_a). The seasonal evolutions of both wind speed and T_s-T_a show increasing seasonal trends similar to those of Q_e and Q_s .

The variability of net heat exchange (Q_N) across the sea surface (Fig.5), in general, indicates a net heat gain of about 150 W m 2 in January and rather a strong net heat loss during February. Maximum get heat loss values go up to 600 W m 2 towards the end of

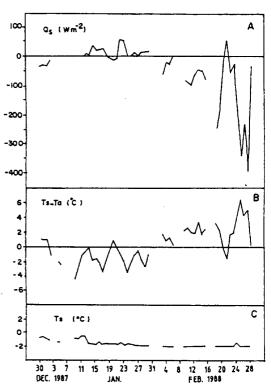


Fig. 4 Daily variation of (A) sensible heat flux, (B) sea-air temperature difference and (C) sea surface temperature.

Frebruary compensating the heat gain of the previous month. On an average, ϱ_{N} during the study period is found to be around 20 W m $^{-2}$

The low net heat gain impedes the faster growth of polynya. The ice cover in the polynya is found to decrease by about 30% only during the observational period. The amount of heat spent for the melting of ice is computed in order to examine the effect of melting of ice on the heat budget of the polynya. A total mass of ice with an average thickness of 1 m and density of 910 Kg m_1^{-3} , works out to be of the order of 5.0×10^{-1} Kg when the entire area of polynya (554 km²) is considered to be covered with pack ice. When 30% of this mass has been melted during the two months observational period, the rate of heat utilised for melting is estimated to be around 17 W m which is very close to the order of net surface heat gain. This balance tries to inhibit a raise in SST it is further expected that the surface heat exchange processes play a minor role in the changes of heat content of upper water column. On the contrary, SST is found to decrease with time. In order to examine the variability of heat content of the water column, a few time series depth-temperature profiles have been examined. From these profiles, it is seen that temporal variations in thermal structure are minimum below 50 m, and accordingly the heat content computations were limited to this depth range. The average rate of heat content change (& H/)t) works out to be around -45 W m

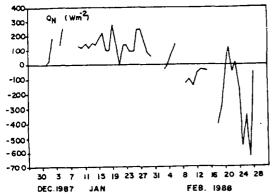


Fig.5 Daily variation of net surface heat flux.

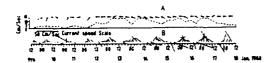


Fig.6 Time series plots of (A) wind speed (----) and wind direction (>>) and (B) stick diagram of currents.

indicating that processes of heat advection from the region are at work.

Necessary computations of heat advection by the currents have been carried out using a limited current meter data set. The required information on the flow velocity has been obtained from an array of current meters moored in the polynya for a limited period in spite of the hostile weather. It is interesting to observe that though the prevailing winds are mainly from east with more or less consistent direction, the variability of currents seems to be dominant with a near diurnal tidal component (Fig.6). This suggests that barotropic processes are important in the coastal circulation of Antarctica.

The vectorially averaged current for the observational period is towards NNE with u and v components of 1.8 and 3.6 cm s respectively. The meredional variation of heat content (3H/3y) across the northern boundary of the polynya is estimated to be 25.4x10 J m . By taking these into consideration, the average advective rate for the observational period works out to be 97 W m . It is interesting to note that this order is comparable with that of change in heat content, further suggesting the importance of advective processes in the heat budget of the shore polynya off Dronning Maud Land Coast, Antarctica. However we point out that a larger time-series data set of currents with an adequate vertical resolution especially in the surface layer is necessary for more reliable estimates of heat advection.

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