### Impact of atmospheric anthropogenic nitrogen on new production in the northern Indian Ocean: Constrained based on satellite aerosol optical depth and particulate nitrogen levels

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### Abstract

Aerosols are one of the significant external sources of soluble reactive nitrogen to the surface ocean and their deposition affects the phytoplankton's primary productivity and associated carbon sequestration. Owing to rapid industrialization over South and Southeast Asia, an increasing trend in atmospheric pollutants was observed over the northern Indian Ocean. To assess the contribution of the aeolian supply of inorganic nitrogen to the northern Indian Ocean (NIO), we have compiled the available data in marine aerosols collected over this basin between 2001 and 2020. Based on the observed relationship of mass load, particulate nitrate and ammonium concentrations over the NIO with the corresponding satellite-derived anthropogenic aerosol optical depth (AAOD) values, the temporal, spatial and long-term variabilities were derived over the NIO for the past two decades. In particular, high aerosol mass load, nitrate and ammonium levels were observed in the coastal aerosols of peninsular India during fall and winter and the least in summer. The basin-wide budget calculations revealed that atmospheric input of inorganic nitrogen to the Arabian Sea is higher (AS; 1.7 Tg/y) compared to the Bay of Bengal (BoB; 0.9 Tg/y) and accounts for ~30% of the total external sources of nitrogen to the NIO. The new production, supported by external sources of nitrogen, contributes to ~23 and 53% of export production to the oxygen minimum zone (OMZ) in the AS and BoB respectively. A significant rate of increase in the aerosol mass load (0.05-1.67  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>), nitrate  $(0.003-0.04 \ \mu g \ m^{-3} \ v^{-1})$  and ammonium  $(0.006-0.11 \ \mu g \ m^{-3} \ v^{-1})$  concentrations was observed between 2001 and 2020, likely because of the growing emission of anthropogenic pollutants over South and Southeast Asia and their subsequent long-range atmospheric transport to the NIO. Overall, these results suggest that enhanced contribution of atmospheric nitrogen may potentially increase 1) the N/P ratio of the surface ocean, which has an impact on the phytoplankton composition, 2) enhanced export production may intensify OMZ and 3) increase in sequestration of atmospheric CO<sub>2</sub>. The decrease in primary production due to global warming is reported due to a decrease in vertical nutrient supply, however, the increase in atmospheric deposition of nutrients may compensate. Therefore, ocean models must couple with atmospheric models for a better forecast of oceanic response to climate change in the NIO.

Keywords: Aerosol optical depth, nutrients, dry deposition, primary production, northern Indian Ocean

### **1. Introduction**

The aeolian transport, from the continent to the ocean, is recognized as one of the important pathways for nutrient supply to the surface ocean<sup>1-4</sup>. An increase in phytoplankton biomass and primary production have been reported due to the deposition of atmospheric nutrients in the global ocean<sup>5-9</sup>. The impact is relatively more significant in the northern hemisphere compared to its southern counterpart mainly due to the larger landmass and more human activity in the former region<sup>10,11</sup>. The global aeolian dissolved inorganic nitrogen (DIN) is reported to be almost equivalent to that of riverine sources<sup>7,12</sup>. Therefore, the long-range transport of atmospheric DIN in the continental outflows to the marginal seas and downwind open ocean waters may have a significant impact on marine primary productivity.

The northern Indian Ocean is surrounded by land from three sides, where one-third of the world population is dwelling close to the coast. Several fast-growing cities are situated along the coast of the northern Indian Ocean resulting in a high contribution of anthropogenic aerosols to the atmosphere<sup>13-</sup> <sup>19</sup>. Hence, this region is classified as one of the hotspot regions in the global ocean with the highest rate of increase in anthropogenic aerosols<sup>20,21</sup>. Seasonal reversal of wind pattern<sup>22</sup> (Scott and McCreary, 2001) leads to variability in sources and composition of aerosols over the north Indian Ocean (NIO)<sup>14,23-25</sup>. Nitrogen is reported as one of the severe limiting nutrients in the NIO (Valsaraj and Rao, 1994; Anju et al., 2020) resulting in the fixation of atmospheric nitrogen (eg: Kumar et al., 2017; Ahmed et al., 2017; Sarma et al., 2020; Kumari et al., 2022a). An increase in primary production and phytoplankton community change due to the deposition of atmospheric nutrients were reported in the Bay of Bengal (Yadav et al., 2016; Kumari et al., 2022b). Based on numerical models and remote sensing data, a significant increase in surface chlorophyll-a associated with atmospheric aerosol deposition has been reported from the NIO<sup>8,26-28</sup>. In addition, efforts have been made to measure the chemical composition of aerosols under the aegis of several national and international programs in India <sup>14,19,23,25,29-34</sup>. Combining several snapshot measurements, Srinivas and Sarin<sup>31</sup> estimated that dry deposition of inorganic nutrients supports ~13% of the surface primary production in the Bay of Bengal (BoB) whereas ~5% in the Arabian Sea (AS) during the dry period. Similarly, Yadav et al.<sup>32</sup> highlighted that ~70% of the total suspended particulate matter (TSP) was deposited within 10 km from the coast and can support 3 to 33% of the coastal surface primary production during fall and winter seasons in the BoB. In addition to primary production, the dry deposition of nitrogen can significantly enhance the N/P ratio in the surface waters which could eventually affect the phytoplankton composition<sup>32</sup>. Most of these estimations were carried out based on snapshot measurements, however, the contribution of atmospheric nutrients to the basin-wide primary/new production is unknown.

Recent studies have found a decrease in the dissolved oxygen concentrations in the sub-surface layers in the BoB and an increase in primary/export production, due to the external supply of nutrients, may turn this region into a dead zone<sup>35</sup>. The weakening of the vertical supply of nutrients and a decline in primary production was reported due to the rapid warming of the tropical Indian Ocean<sup>36,37</sup>. On the other hand, the increasing rate of atmospheric pollutants over South and Southeast Asia was reported with a dominant increase in NO<sub>x</sub> and SO<sub>x</sub> from the middle of  $2000^{38}$ . Such a rapid increase in NO<sub>x</sub> may enhance particulate nitrogen deposition over the northern Indian Ocean leading to a possible increase in primary production in future. It is also possible that rapid warming associated with a weaker vertical supply of nutrients (upwelling) may be compensated by the deposition of nutrients from the atmosphere to sustain primary production. Sporadic studies on the chemical composition of aerosols revealed a significant fraction of aerosols over the northern Indian Ocean were contributed by anthropogenic sources<sup>14,19,23,25,31,32,34</sup>. The increase in anthropogenic aerosol optical depth (AAOD) was reported in the BoB over the past two decades<sup>21</sup>. However, the trends in the concentrations of reactive nitrogen and their influence on surface primary production due to their subsequent deposition on the surface ocean are unknown. Therefore, the objective of the present study is to 1) quantify the spatial, temporal and long-term pattern of inorganic nitrogen concentrations in the marine boundary layer (MBL) of the northern Indian Ocean using the available measured data as well as satellite-derived AOD over the northern Indian Ocean, 2) quantify inorganic nitrogen inputs through atmospheric deposition and their contribution to new production and 3) evaluate the long-term trends in atmospheric deposition of inorganic nitrogen over the northern Indian Ocean.

### 2. Methodology and study area

### 2.1. Mass load, nitrate and ammonium concentrations in the marine aerosols over the NIO

The measured mass load, nitrate and ammonium concentrations of dry aerosols over the NIO between 2006 and 2020 data were used in this study during different seasons. The coverage was relatively good in the northeastern Indian Ocean (BoB) compared to the northwestern Indian Ocean (AS). The dates of the collection of samples and the region of the collection were given in Table 1. Ambient aerosol mass load over the NIO was assessed by the difference in filter weights before and after collection. The concentrations of nitrate and ammonium were measured using the Ion Chromatography techniques with suppressed conductivity detection. The details of the measurements and their precisions were given elsewhere<sup>14,23,25,32</sup>.

### 2.2. Computation of anthropogenic aerosol optical depth (AAOD)

Since inorganic nitrogen (nitrate and ammonium) mainly enters the atmosphere through anthropogenic sources, the fraction of AOD contributed by the anthropogenic sources needs to be computed from the total AOD. We have obtained total AOD data at 1 x 1° grid resolution at 470 nm, 550 nm and 860 nm wavelengths from the Moderate Resolution Imaging Spectroradiometer Terra satellite (MODIS) collection 6.1 and level 3 from 2001 to 2019 (https://ladsweb.modaps.eosdis.nasa.gov/search). The Level 3 (L3) MODIS AOD atmosphere products data is a global-gridded science product. The fine mode fraction of AOD was computed from spectral AOD and fine-mode AOD (given as aerosol optical Depth small-Ocean)<sup>39</sup>. The error on the total AOD is  $\pm 0.03+5\%$  of AOD<sup>40</sup>. The fraction of anthropogenic AOD (AAOD) from total AOD was estimated following Kaufman et.al.<sup>41</sup>. The estimation of the anthropogenic component from MODIS total AOD uses several assumptions. They are a) the fine fraction of aerosol is constant to the AOD for a given aerosol type. Smoke and pollution are assumed to be mostly contributed to fine aerosols, whereas dust and marine components contribute to coarse aerosols. b) All smoke is considered to be of anthropogenic origin and all dust originated naturally. c) It is assumed that wildfires contribute ~20% of biomass burning and anthropogenic sources contribute up to 10% of the dust. d) MODIS-derived fine fraction is highly correlated with a true fine fraction with small errors. e) Based on AERONET and MODIS analysis<sup>41,42</sup> the baseline marine aerosol optical thickness is assumed to be  $0.06 \pm 0.01$  under calm conditions.

According to Kaufman et al.<sup>41</sup> total aerosol optical thickness ( $\tau_{550}$ ) is defined as the sum of anthropogenic ( $\tau_{anth}$ ) from pollution and smoke + dust ( $\tau_{dust}$ ) + baseline marine ( $\tau_{mar}$ ) sources. The same can be written as

 $\tau_{f} = f_{550} \ x \ \tau_{550} = f_{anth} \ x \ \tau_{anth} + f_{dust} \ x \ \tau_{dust} + f_{mar} \ x \ \tau_{mar} \quad (1)$ 

Here  $f_{mar}$ ,  $f_{dust}$ , and  $f_{anth}$  are fraction fine particles from marine, dust and anthropogenic sources and they are derived from MODIS whereas  $\tau_{mar}$  is determined from AERONET. The total fine fraction ( $f_{550}$ ) and optical thickness ( $\tau_{550}$ ) were derived by MODIS. Based on this,

$$\tau_{anth} = [(f_{550} - f_{dust}) \tau_{550} - (f_{mar} - f_{dust}) \tau_{mar}]/(f_{anth} - f_{dust}) \quad (2)$$

The complete details of the estimations and constants used are given in Kaufman et al.<sup>41</sup>. In the present study, we used MODIS collection 6.1 AOD data whereas Kaufman et al.<sup>41</sup> used collection 4 data to derive the constants, therefore, we calculated the constants  $f_{mar}$ ,  $f_{dust}$ , and  $f_{anth}$  using MODIS collection 6.1 data for the same regions and periods that of Kaufman et al.<sup>41</sup>. We obtained  $f_{mar}$ ,  $f_{dust}$ , and  $f_{anth}$  as

0.32, 0.48, and 0.77 respectively and the computation of the fraction of anthropogenic aerosols carries an uncertainty of up to  $\pm 30\%$ .

The weighted linear regression analysis was conducted on the monthly mean data of anthropogenic AOD, aerosol nitrate and ammonium concentrations in the selected regions in the northwest AS (NWAS), northeast AS (NEAS), southwest AS (SWAS), southeast AS (SEAS), northern BoB (NBoB), southeast BoB (SEBoB) and southwest BoB (SWBoB; Figure 1) regions between 2001 and 2020 to obtain long-term trends. The NEAS and NBoB boxes represent the regions that receive high anthropogenic aerosols from Indo-Gangetic Plain (IGP) region, whereas the SEAS region receives from mainland India. Both NWAS, SWAS in the AS and SWBoB, SEBoB boxes in the BoB are away from the anthropogenic aerosol source regions.

### 2.3. Dry depositional flux of nitrate and ammonium

The dry depositional fluxes of NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup> were computed following:

Dry-Deposition flux =  $C_{measured} * V_{dry} --- (3)$ 

Where  $C_{measured}$  is the concentration of nitrate or ammonia in the aerosols and  $V_{dry}$  is the dry deposition velocity, which is estimated to be 1.0 and 0.1 cm sec<sup>-1</sup>, respectively for nitrate and ammonia in the  $PM_{10}^{-1}$ . The low  $V_{dry}$  for  $NH_{4^+}$  (0.1 cm sec<sup>-1</sup>) is considered due to its predominant occurrence in the fine mode<sup>23,43,44</sup>. The uncertainty associated with the computation of dry deposition fluxes can be by a factor of 3 due to several factors including wind speed, terrain roughness, particle composition etc<sup>31</sup>.

### 2.4. Computation of long-term trends in the selected region

To understand the long-term changes of all these parameters on a seasonal scale, data are prepared into seasonal means (winter-December, January, February; Spring – March, April and May; Summer – June, July, August and September and Fall – October and November) and then liner trend analysis was performed on all the parameters.

We have determined the trends and statistical significance of the parameters studied following Santer et al<sup>45</sup>. The trends of the parameters were computed using a least-square fit method for each seasonal data as:

x(t) = a + bt, for t = 1, ..., nt

The significance of the computed trends was evaluated following Santer et al.<sup>45</sup>, in which regression residuals (e(t)) are defined as:

 $e(t) = x(t) - x_m(t)$ ; fort = 1, ..., nt

 $x_m$  is the mean value and for statistically independent values of e(t), the standard error of the slope, b, is defined as:

$$Sb = Se/(\sum (t-t_m)^2)^{0.5}$$

where  $t_m$  is the mean and  $Se^2$  is the variance of the residuals about the regression line, is given by

$$Se^2 = (1/n_t-2)^* \sum e(t)^2$$

Whether the e(t) trend is significantly different from zero is tested by computing the ratio between the estimated trend and its standard error

### $t_b = b/s_b$

We noticed that e(t) is not statistically dependent hence autocorrelation method involving sample size is used following Santer et al<sup>45</sup>. The seasonal trend analysis results and their significance between 2001 and 2020 are given in Table 2.

### 3. Results

### 3.1. Importance and meteorological condition of the Northern Indian Ocean (NIO)

The NIO is surrounded by the deserts in the Oman, Somalia, Iran and Thar regions in the west (Arabian Sea), and surrounded by highly populated underdeveloped nations such as Bangladesh, Burma, Thailand etc., in the eastern basin (Bay of Bengal). The NIO is influenced by the seasonally reversing monsoonal winds that blow from the northeast direction during winter (November to February, called northeast monsoon) and southwest direction during summer (June to September; called southwest monsoon; Wiggert et al., 2005). The air mass over the NIO is strongly influenced by the continental pollution sources (eg: biomass burning and fossil-fuel combustion) in winter (Bikkina et al., 2011; 2021; Kumar et al., 2012; Ashwini et al., 2020). The Arabian Sea receives mineral dust particles in summer from the dessert region (Tindale and Pease, 1999; Prospero et al., 2002). In the case of the Bay of Bengal, the continental pollution from the Indian subcontinent is received during summer and Southeast Asia during winter (Kumari et al., 2022). More than 80% of the annual precipitation occurs during summer and scarce or episodic rainfall occurs during winter (Soman and Kumar, 1990).

### 3.2. Distribution of anthropogenic aerosols (AAOD) over the northern Indian Ocean

The climatological mean AAOD (2001-2020) displayed significant spatial and temporal variations over the northern Indian Ocean, which may be partially attributed to meteorological factors such as the seasonal reversing of wind patterns. The levels of AAOD were relatively higher along the Indian coast and rapidly decreased towards the open oceanic region (Fig. 1). The levels of AAOD increased

rapidly from September to April and lower levels were noticed between May and August (Fig. 1). The high levels of AAOD were observed along the east coast of India between November and March with the peak in December and January. In contrast, high levels of AAOD were observed from October to March with a peak in October to November along the west coast of India. A clear north-south gradient in AAOD levels was observed in the BoB with relatively higher values in the north and a decreasing trend towards the south. Such decreasing pattern was absent in AAOD and was almost missing in the AS.

# 3.3. Distribution of measured mass load, nitrate and ammonium concentrations in aerosols over the northern Indian Ocean

The distribution of measured climatological mass load, nitrate and ammonium concentrations in the aerosols were presented in Figure 2a. The mass load was higher along the Indian coast and northern BoB compared to the offshore and southern BoB. Higher concentrations of nitrate and ammonium were observed along the Indian coast than offshore (Fig. 2b,c). AAOD values displayed a significant linear relationship with the mass load, nitrate and ammonium concentrations in the AS and BoB (Fig. 3). The slope of the relationship was different in the two basins. For instance, the slope of the relationship of AAOD with mass load, nitrate and ammonium was almost double in the BoB (115.8, 3.4, and 7.6 respectively) than in the AS (54.8, 1.7 and 3.3 respectively). The relationship yielded the following regression equations:

 $Mass = 115.84 * AAOD + 11.322: r^{2}=0.54 (p<0.001) (BoB) --- (4)$  $NO_{3}^{-} = 3.42 * AAOD + 0.004: r^{2}=0.51 (p<0.001) (BoB) --- (5)$  $NH_{4}^{+} = 7.65 * AAOD - 0.59: r^{2}=0.58 (p<0.001) (BoB) --- (6)$  $Mass=54.85 * AAOD + 21.47: r^{2}=0.62 (p<0.001) (AS) --- (7)$  $NO_{3}^{-} = 1.71 * AAOD + 0.31: r^{2}=0.69 (p<0.001) (AS) --- (8)$ 

$$NH_4^+ = 3.29 * AAOD + 0.32: r^2 = 0.68 (p < 0.001)$$
 (AS) --- (9)

3.4. Temporal and spatial variations in the mass load, nitrate and ammonium concentrations in the aerosol over the northern Indian Ocean

The monthly variations in climatology (2001-2020) of the computed mass load, nitrate and ammonium concentration varied from 0.5 to 100, 0 to 2 and 0 to 3.5  $\mu$ g m<sup>-3</sup> respectively in the NIO. Higher concentrations of mass, nitrate and ammonium concentrations were observed in the BoB than in the AS. The high concentrations of mass (>50  $\mu$ g m<sup>-3</sup>), nitrate (>1.5  $\mu$ g m<sup>-3</sup>) and ammonium (>2.5  $\mu$ g m<sup>-3</sup>) were observed along the coastal BoB with high concentrations between October and April than in

other seasons (Fig. 4-6). On the other hand, high concentrations of nitrate and ammonium were observed in the south-eastern AS from October to April compared to other months (Fig. 5 and 6). A perennial low concentration of nitrate (<0.5  $\mu$ g m<sup>-3</sup>) and under detection limits of ammonium was observed in the southern BoB. Relatively high concentrations of nitrate and ammonium were observed in the offshore regions in the AS from October to January while it was from January to April in the case of the BoB.

### 3.5. Dry depositional flux of nitrate and ammonium in the northern Indian Ocean

The dry depositional flux of nitrate and ammonium varied between 0-1.8 and 0-0.3 mgN m<sup>-2</sup> d<sup>-1</sup> in the northern Indian Ocean respectively (Fig. 7-8). Both nitrate and ammonium fluxes were higher in the north-western BoB (>1.4 and >0.25 mgN m<sup>-2</sup> d<sup>-1</sup> respectively) followed by south-eastern AS from October to April (0.6-1.0 and 0.1-0.15 mgN m<sup>-2</sup> d<sup>-1</sup> respectively). Assuming the C/N ratios of 6.6, and deposited nitrate and ammonium were completely utilized by the phytoplankton, the contribution of atmospheric deposition of inorganic nitrogen to the primary production in the surface ocean was 0-18.5 mgC m<sup>-2</sup> d<sup>-1</sup> in the northern Indian Ocean (Fig. 9).

### 3.6. Long-term trends in AAOD, mass, nitrate and ammonium fluxes in the northern Indian Ocean

The seasonal long-term trends in AAOD, mass, nitrate and ammonium fluxes were derived in the selected regions (Fig. 1) in the northern Indian Ocean. The rate of AAOD was significant in all the regions at a 99% confidence interval, except in the SWBoB and SEBoB where it is either low or insignificant (Table 2). The rate of increase in AAOD varied between 0.002 and 0.014/y with lower rates in the SWAS (0.003-0.005/y) and higher in the NEAS, SEAS and NBoB (0.007-0.014/y). The lowest rate of increase with low significance (p<0.05) was observed in the SWBoB (0.002-0.003/y). Within the seasons, a high rate of increase in AAOD was noticed during the fall monsoon and either decreasing or insignificant trends were observed during summer.

The rate of increase in mass load varied between 0.11 and 1.67  $\mu$ g m<sup>-3</sup> y<sup>-1</sup> in the northern Indian Ocean. The highest rate of the mass load was observed in the NBoB (0.59 to 1.67  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>) during all seasons, except summer, followed by SEAS (0.66-0.80  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>) and NEAS (0.12 to 0.66  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>) and the least rate of increase was noticed in the SWAS (0.14-0.26  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>) and SEBoB (-0.04 to 0.25  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>). Except in the NBoB and SEAS, the rate of mass load decreased during summer (Table 2).

The rate of increase in nitrate and ammonium in the aerosols ranged between 0.003 to 0.049  $\mu$ g m<sup>-3</sup> y<sup>-1</sup> and 0.01 to 0.11  $\mu$ g m<sup>-3</sup> y<sup>-1</sup> respectively in the northern Indian Ocean. The rate of increase in nitrate was higher in the SEAS and NBoB (0.02 to 0.05  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>) during all seasons, except summer. Except

in NBoB and SEAS, negative (decrease) trends in nitrate were noticed during summer. A higher rate of increase in ammonium was observed in the NBoB and SEAS (0.04 to 0.11  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>) than in other regions (Table 2).

### 4. Discussion

# 4.1. Spatial variations in atmospheric deposition of nitrogen and its contribution to surface primary production in the northern Indian Ocean

The northern Indian Ocean receives anthropogenic rich aerosols during winter associated with northeasterly winds while the dominance of mineral dust and sea salt is observed during summer months when prevailing winds are southwesterly<sup>14,16,46-48</sup>. The primary source of dust to the northern Indian Ocean comes from the Nubian desert, the Arabian Peninsula, Iran, Pakistan, Afghanistan and North-West India with maximum activity between May and October<sup>49-52</sup>. The fraction of AAOD to the total AOD was higher during winter and spring (November-April; 60-70%) than in summer (30-40%) over the northern Indian Ocean due to the influx of anthropogenic rich aerosols derived from the surrounding continents where >30% of the global population dwells<sup>21,53</sup>. The prennial high levels of AAOD were observed along the Indian coast with the rate of increase being the highest over the globe<sup>20,21</sup> and it was attributed to a significant increase in the contribution of anthropogenic emissions from the Indo-Gangetic Plain (IGP) and southeast Asia.

The climatological mean (2001-2020) derived mass load, nitrate and ammonium concentration were high along the Indian coasts from October to April due to prevailing terrestrial winds resulting in the input of anthropogenic aerosols. The sharp gradient in aerosol mass, nitrate and ammonium concentration was observed from the coast towards the open oceanic region mainly due to the proximity of the emission source region. During a similar period, the western BoB receives aerosols from the biomass burning from the Indo-Gangetic Plain (IGP) region and southeast Asia due to dominant winds from the northeast direction resulting in enhanced mass load, nitrate and ammonium concentration along with the coastal and northern BoB<sup>54,55</sup>. These are the least during summer due to washout from the atmosphere via heavy monsoonal precipitation<sup>56</sup>.

The AAOD displayed a linear relationship with nitrate and ammonium concentrations in the aerosols at 99% significance levels (See equations 4-9) suggesting that anthropogenic sources are dominant contributors of dissolved inorganic nitrogen to the aerosols over the northern Indian Ocean. Perennial high concentrations of nitrate and ammonium were observed along the Indian coast (Fig. 5,6); particularly high levels from October to April months. The dry depositional flux of nitrate and ammonium varied from 0-1.8 and 0-0.3 mgN m<sup>-2</sup> d<sup>-1</sup> in the northern Indian Ocean with higher fluxes

along the Indian coastal region and northern BoB due to proximity to the source emissions (Fig. 7,8). The deposition of atmospheric nitrogen may enhance primary production in the surface ocean<sup>1,31</sup>. Recently Kumari et al. (2022) measured an increase in primary production by 3-19% due to the deposition of atmospheric aerosols in the coastal Bay of Bengal whereas Yadav et al. (2016) noticed increase by 3 to 33%. Episodic increase in phytoplankton blooms associated with dust storms was reported in the northern Indian Ocean and attributed to the deposition of atmospheric nutrients<sup>8,26-28,51</sup>. Severe nitrogen limitation to primary production was reported in the northern Indian Ocean (Anju et al., 2020; Sarma et al., 2013; Valsaraj and Rao, 1994). Assuming that the deposited inorganic nitrogen (nitrate and ammonium) on the surface ocean is completely utilized by the phytoplankton, it may support the primary production of 0-18.5 mgC m<sup>-2</sup> d<sup>-1</sup> in the northern Indian Ocean (Fig. 9b, c).

To examine the contribution of primary production supported by atmospheric deposition of inorganic nitrogen to total primary production, the surface primary production data was measured using the radio-active carbon (14C) tracer incubation method for comparison. The seasonal and spatial variability in primary production in the northern Indian Ocean was measured between 1994 and 2019 during various national and international programs<sup>57-63</sup>. The upper 5 m integrated primary production varied between 7.8 and 387 mgC m<sup>-2</sup> d<sup>-1</sup> in the AS and 6.8 to 225 in the BoB (Fig. 9a). Assuming the entire dry deposition of inorganic nitrogen is utilized by the phytoplankton, the air-to-sea flux of nitrogen is converted to equivalent carbon produced by the phytoplankton primary production using a canonical Redfield ratio (C/N ratio of 6.6). The primary production supported by atmospheric borne nitrate and ammonium between 2001 and 2020 was used to evaluate its contribution to the total primary production in the upper 5 m depth. In the case of primary production measurements made before the year 2000, the primary production derived from the climatological mean nitrate and ammonium deposition rates was considered for comparison due to the lack of AOD data. The contribution of primary production, supported by atmospheric deposition of nitrogen, to total primary production in the upper 5 m of the water column ranged between 1.1 and 45% (mean of 14±7%) in the BoB whereas it was 1 to 54 (11±8%) in the AS (Fig. 9d). Based on the depositional flux estimated using measured dissolved inorganic nitrogen, the contribution of atmospheric nitrogen to surface primary production was reported to vary between 3 and 33% during different seasons with a maximum during winter in the coastal  $BoB^{23,32}$ . Srinivas et al.<sup>23</sup> estimated that <13% of the primary production in the entire BoB is supported by atmospheric deposition of nitrogen whereas it is <5% in the AS<sup>31,64</sup>. The present estimations suggested that atmospheric deposition of nitrogen contributed almost equally to the surface primary production in the AS and BoB than hitherto estimated that higher contribution in the latter than the former  $basin^{23,31}$  (Fig. 9d).

### 4.2. Contribution of external sources of nitrogen to the northern Indian Ocean

The northern Indian Ocean is a highly productive region within the Indian Ocean with relatively higher production in the AS than BoB due to various oceanic processes which include the occurrence of strong upwelling and convective mixing during summer and winter respectively<sup>58,59,63</sup>. In contrast, the absence of such strong mixing in the BoB due to strong salinity stratification results in low primary production<sup>59,65</sup>. Due to the proximity to the land, the northern Indian Ocean receives a significant amount of nutrients through land-ocean (rivers and groundwater) and ocean-atmosphere (atmospheric deposition) boundaries. Due to stratification during non-monsoon periods with low inorganic nitrogen inputs to the surface, nitrogen fixation occurs and contributes significantly to the primary production in the northern Indian Ocean. Based on the available data in the recent literature and present depositional flux of nutrients from the atmosphere estimates, the contribution of external sources of nitrogen to the northern Indian Ocean was computed.

Several rivers, such as Ganges, Brahmaputra, Irrawady-Salween system, Godavari, Krishna, Mahanadi etc, opened to the BoB, whereas Narmada, Tapti, Mandovi-Zuari etc rivers are opened to the AS<sup>66</sup>. The river discharge to the BoB is 3 times higher than that of the AS due to the discharge of more rivers in the former basin<sup>67</sup>. Krishna et al.<sup>66</sup> estimated riverine inorganic nitrogen inputs to the AS and BoB to be  $0.10\pm0.03$  and  $1.74\pm0.43$  TgN/y and contribute to  $0.7\pm0.2$  and  $11.5\pm3$  TgC/y of new production respectively based on C/N ratio of 6.6 (Table 3).

The northern Indian Ocean has a very long coastline and the exchange of nutrients through submarine groundwater discharge (SGD) is a possible source of nutrients. The higher concentrations of nutrients, by an order of magnitude, in the groundwater than in coastal waters were reported<sup>51</sup>. The estimated SGD flux of dissolved inorganic nitrogen was estimated to be  $0.06\pm0.03$  and  $0.102\pm0.02$  TgN/y to the AS and BoB, which supports  $0.4\pm0.2$  and  $0.7\pm0.1$  TgC/y of new production respectively.

The estimated long-term mean (2001-2020) flux of atmospheric nitrate and ammonium varied between 0-1.8 and 0-0.3 mg N m<sup>-2</sup> d<sup>-1</sup> in the northern Indian Ocean respectively with higher rates in the western BoB and eastern AS. The estimated deposition of atmospheric inorganic nitrogen over the entire AS and BoB is estimated to be  $1.7\pm0.4$  and  $0.9\pm0.2$  TgN/y respectively and it would support 11.2 and ~6.0 TgN/y of new production respectively in the AS and BoB.

The occurrence of nitrogen fixation was reported both in the AS and BoB during the inter-monsoon period. In the case of the AS, most of the measurements were taken in the eastern basin and few measurements were carried out in the western basin. Nitrogen fixation rates varied between 124 and 1500  $\mu$ mol N m<sup>-2</sup> d<sup>-1</sup> in the AS<sup>68-72</sup> with a mean of 532±547  $\mu$ mol m<sup>-2</sup> d<sup>-1</sup>, excluding the one outlier of

36000  $\mu$ mol N m<sup>-2</sup> d<sup>-1</sup>, and extrapolating to the entire basin during spring inter-monsoon amounts to 4.4 TgN y<sup>-1</sup> in the AS. The N<sub>2</sub> fixation rates in the BoB varied between 0.7 and 75.2  $\mu$ mol m<sup>-2</sup> d<sup>-1</sup> <sup>73-75</sup> with a mean of 10.5±12  $\mu$ mol m<sup>-2</sup> d<sup>-1</sup> and extrapolating to the entire BoB amounts to 0.2±0.1 TgN y<sup>-1</sup>. The equivalent to new production is estimated based on a C/N ratio of 6.6. Based on the synthesis of measured N<sub>2</sub> fixation data, the contribution of new production due to N<sub>2</sub> fixation is estimated as 29.0 and 1.3 TgC/y in the AS and BoB respectively. The low N<sub>2</sub> fixation in the BoB is attributed to the removal of phosphate in association with suspended particulate matter<sup>76</sup>.

The total external sources of inorganic nitrogen to the AS and BoB come to  $6.3\pm4$  and  $2.0\pm0.4$  TgN/y respectively. The contribution of nitrogen fixation dominates (70%) followed by atmospheric deposition (27%) in the AS whereas river discharge (60%) followed by atmospheric deposition (30%) contributed to total external sources in the BoB. The contribution of SGD is minimal (1-3%) in both AS and BoB (Table 3). The external sources support  $41.3\pm26$  and  $19.5\pm3$  TgC of new (primary) production in the AS and BoB respectively.

Since most of the new production is exported to the twilight zone<sup>77</sup> (Dugdale and Goering, 1967), the contribution of new production by external sources to the export production is compared. The export production at 100 m was measured in the northern Indian Ocean from 1994 to 1996 and 2013 to 2015<sup>78-80</sup>. Seasonal variability in export production was measured in the AS whereas only measurements were conducted during the non-monsoon season in the BoB, therefore, some seasonal bias is anticipated in the latter basin. Based on these estimates, the export flux to the twilight zone was accounted to be 182 and 36 TgC/y in the AS and BoB respectively<sup>78-80</sup>. The total external sources of nitrogen supported new production contributed to 23 and 53% of the export production in the AS and BoB respectively and the remaining export production is contributed by nutrient upwelled from the subsurface water, either wind-driven coastal upwelling (Morrison et al., 1998) or eddy-driven upwelling (Sarma et al., 2018) and regenerated nutrients through microbial processes. Among the external sources, nitrogen fixation (16%) and atmospheric deposition (6%) contributed significantly to the export production in the AS whereas river discharge (31%) and atmospheric deposition (17%) contributed to the BoB. Therefore, the contribution of atmospheric deposition is significant not only to surface production but also to export production to the twilight zone in the northern Indian Ocean.

In this study, we have estimated the contribution of inorganic nitrogen to the surface ocean through dry depositional flux. Jickells et al. (2017) reported that organic nitrogen contributes about 1/3 of the total nitrogen deposition. A dominant fraction of organic nitrogen (40-50%) to the total nitrogen was reported in the Arabian Sea (Bikkina et al., 2021) and the Bay of Bengal (Bikkina et al., 2011).

Recently Rao et al. (2021) reported that dissolved organic nitrogen contributes to ~50% of the primary production in the Bay of Bengal in the absence of inorganic nitrogen. Due to the lack of spatial data on dissolved organic nitrogen in the aerosols, we could not account for their contribution to the total primary production in this study. On the other hand, the NIO is highly influenced by monsoonal precipitation and the contribution of wet deposition of nitrogen may be a highly significant nitrogen source to the surface ocean. Baker et al. (2021) reported that the acidity of the aerosols changes the lifetime of nitrate in the atmosphere and increases the ammonium resulting in the long-range transport of ammonium and deposition close to the coast in the case of nitrate under acidic conditions. The acidic nature of the aerosols over the Bay of Bengal was reported recently (Kumari et al., 2022) and its impact on portioning of nitrate and ammonium is unclear. Nevertheless, the role of organic nitrogen and wet deposition to total nitrogen deposition and its contribution to the external sources and surface primary production needs further studies.

Bristow et al.<sup>35</sup> found that dissolved oxygen in the subsurface layers is at the tipping point with reference to oxygen levels in the oxygen minimum zone (OMZ) and an increase in external sources of nutrients may enhance primary and export productions that may further intensify oxygen levels leading to the formation of dead zones. In this context, our study highlights the contribution of atmospheric sources is significant, especially in the BoB (17%) than in the Arabian Sea (6%), which may intensify OMZ formation. Kurokawa and Ohara<sup>38</sup> found a rapid rate of increase in NOx and SOx in the atmosphere over South and Southeast Asia since the middle of 2000. This may enhance nitrogen deposition over the northern Indian Ocean which may increase primary and export production. These results underscore the need to examine the long-term trends of atmospheric nutrient concentrations in synchronous with the dissolved oxygen levels in the subsurface waters of both basins.

### 4.3. Long-term trends in atmospheric deposition of dust, nitrogen and its impact on primary production

The long-term trends in the AAOD, mass load, nitrate and ammonium concentrations in the selected regions were shown in Table 2. The rate of increase in AAOD was more than double in the eastern basin (NEAS and SEAS)  $(0.007\pm0.002/y)$  than western basin (SWAS and NWAS)  $(0.004\pm0.001/y)$  due to proximity to the source region of the Indian subcontinent in the former regions, except summer. The rate of increase in AAOD in the northern BoB was the highest (0.01/y) and it was insignificant in the southern BoB. The rate of increase in AAOD in the BoB was higher than earlier reports between 2000 and 2009  $(0.0077\pm0.0003/y)$  but lower in the case of the Arabian Sea  $(0.006\pm0.0002/y)^{20.21}$  observed that the rate of increase in AAOD was significant between 2001 and 2019 in the northwestern and western coastal BoB but insignificant in the other regions (Yadav et al., 2021).

The computed mass load, nitrate, and ammonium concentrations in the aerosols displayed an increasing trend in the SWAS during all seasons but displayed negative trends in summer. A relatively higher rate of increase in mass load was observed in the eastern basin (NEAS and SEAS) than in the western basin (SWAS and NWAS) due to proximity to the source region from the Indian subcontinent. Within the seasons, a higher rate of increase in mass load was observed during fall monsoon in the NEAS, SWAS and NWAS whereas winter and spring in SEAS than in other seasons (Table 2). The annual mean rate of increase in mass concentration was also higher in the NEAS (0.221±0.182 µg m<sup>-</sup>  $^{3}$  y<sup>-1</sup>), SEAS (0.443 µg m<sup>-3</sup> y<sup>-1</sup>) than in SWAS (0.103±---µg m<sup>-3</sup> y<sup>-1</sup>) and NWAS (0.109±---µg m<sup>-3</sup> y<sup>-1</sup>) suggesting that the eastern basin is receiving a significant amount of anthropogenic aerosols since past two decades than a western basin of AS. The highest increase in mass load was observed in the NBoB  $(1.18 \ \mu g \ m^{-3} \ y^{-1})$  due to the receipt of major sources from the IGP and Southeast Asia<sup>21</sup>. The rate of increase in mass load was the least in the SEBoB (0.05  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>) and SWBoB (0.34  $\mu$ g m<sup>-3</sup> y<sup>-1</sup>) due to being far from the source region. This study suggests that a significant increase in the mass of aerosols was noticed in the eastern AS and northern BoB in the northern Indian Ocean which may decrease the rate of increase in SST. Dong et al.<sup>81</sup> observed that the rate of warming decreased in the northern Indian Ocean due to AAOD in the atmosphere. This impact was found to be stronger over the BoB than the AS due to higher levels of AAOD over the former than in the latter basin. Sarma et al.<sup>82</sup> also reported a decrease in the rate of warming in the north-western BoB due to an increase in atmospheric pollutants between 1991 and 2011. However, the role of AAOD on SST needs further evaluation in the northern Indian Ocean.

Significant increasing annual mean trends in nitrate and ammonium concentrations were observed in the eastern (0.007-0.013 and 0.012-0.023  $\mu$ g m<sup>-3</sup> y<sup>-1</sup> respectively) than western AS (0.002-0.003 and 0.004-0.005  $\mu$ g m<sup>-3</sup> y<sup>-1</sup> respectively; Table 2) suggesting that the former region of the basin receiving higher nutrients through dry deposition. The increasing annual mean trends in nitrate and ammonium were higher in NBoB (0.034 and 0.076  $\mu$ g m<sup>-3</sup> y<sup>-1</sup> respectively) than in SEBoB and SWBoB (0.001-0.01 and 0.005-0.013  $\mu$ g m<sup>-3</sup> y<sup>-1</sup> respectively). The depositional rate of inorganic nitrogen concentrations increase by 135% in the NBoB followed by SWBoB (34%) and the least in the SEBoB (13%) between 2001 and 2020. In the case of AS, the increase in inorganic nitrogen was more in the NEAS (30%) and SEAS (44%) compared to NWAS (14%) and SWAS (11%). The mean rate of increase in primary production between 2001 and 2020 due to deposition of atmospheric inorganic nitrogen was higher in the NBoB (2.5 gC m<sup>-2</sup> y<sup>-1</sup>) followed by eastern AS (1.2-1.8 gC m<sup>-2</sup> y<sup>-1</sup>), western AS (1.0-1.2 gC m<sup>-2</sup> y<sup>-1</sup>) and southern BoB (0.6-0.8 gC m<sup>-2</sup> y<sup>-1</sup>). Bindoff et al.<sup>83</sup> found that a decline in primary production is projected by 50 gC m<sup>-2</sup> y<sup>-1</sup> in the western AS whereas 30 gC m<sup>-2</sup> y<sup>-1</sup> in the eastern

AS and southern BoB, in contrast, an increase by 20 gC m<sup>-2</sup> y<sup>-1</sup> was reported in the NBoB and SEBoB between 2081-2100 relative to 1950-1900 under RCP 8.5 scenario. The decline in primary production by end of this century was attributed to the warming of the ocean, an increase in stratification and a weakening of nutrient supply through vertical mixing. In contrast, our study suggests that enhanced atmospheric pollution in future may compensate for nutrient decrease due to warming and stratification that may decrease the projected declining rates. Since the rapid increase in atmospheric pollution is already reported in South and South Asia<sup>38</sup> over the past two decades, atmospheric nutrient contribution may be significant in future and needs further evaluation.

### **Summary and Conclusion**

The dry deposition of atmospheric nitrogen is hypothesized to be an important source of nutrients to the northern Indian Ocean. To quantify this, temporal and spatial variations in concentrations and depositional rates of inorganic nitrogen in the atmosphere were computed based on the relationship between measured nitrogen in the aerosols and satellite-derived anthropogenic aerosol optical depth (AAOD). The perennial high concentrations of mass load, nitrate and ammonium were observed in the aerosols over the eastern AS and western BoB with characteristic seasonality. A significant decrease in AAOD levels and concentrations of nitrogen was observed in the open oceanic region. This suggests that the Indian subcontinent could be one of the potential sources of pollutants, along with emissions from Southeast Asia, to the northern Indian Ocean. Relatively higher concentrations of nitrate and ammonium were observed in the aerosols during fall and winter and the least during summer due to heavy precipitation. The contribution of atmospheric deposition of inorganic nitrogen amounted to 1.7 and 0.9 TgN  $y^{-1}$  to the AS and BoB respectively. It accounts for ~30% of the external sources (river discharge, N<sub>2</sub> fixation and submarine groundwater discharge) of inorganic nitrogen to the northern Indian Ocean. The total external sources contribute 23 and 53% of the export production to the oxygen minimum zone (OMZ) in the AS and BoB respectively. The long-term variability in the mass load, nitrate and ammonium concentrations in the aerosols between 2001 and 2020 suggests a significant rate of increase due to a perennial increase in atmospheric pollution. This may result in a significant increase in primary and export production to OMZ and a decrease in the warming rate due to the blanketing effect.

The northern Indian Ocean harbours one of the most intense OMZs in the world and the oxygen levels in the OMZ are controlled by primary/export production and circulation of the intermediate waters. The nanomolar levels of oxygen are observed in the OMZ in the northern BoB and reported to be at the tipping point<sup>35</sup>. This study suggests that atmospheric deposition is contributing to 30% of external nitrogen input to the northern Indian Ocean with a perennial increase in its contribution, therefore, the

organic matter flux to the OMZ may increase in future. This may lead to further intensification of the OMZ in the northern Indian Ocean. The atmospheric deposition not only increases inorganic nitrogen concentration but also the N/P stoichiometric ratio which may have a significant impact on the phytoplankton composition and therefore food web dynamics in the northern Indian Ocean<sup>32</sup>. While this study synthesizes a snapshot picture of aerosol composition in different seasons, however, to get a robust trend in nutrient supply and assess their impact on surface water processes, the time-series observations of atmospheric aerosol composition and associated nutrients over the northern Indian Ocean are essential, which enable us to understand the role of the human impact on ecosystem structure. Nevertheless, this study suggests that a perennial increase in an atmospheric nutrient deposition may compensate for the decrease in vertical nutrient supply due to stratification under the climate change scenario. Therefore, the ocean models must be coupled with the atmospheric models for a better forecast of a possible change in the ocean's primary production due to climate change.

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Cruise No.	Dates	Area	Reference
AS 430-528	11 April to 7 May 2006	Arabian Sea	14
SSK 60	24 Jan to 1 Feb 2014	Arabian Sea	Unpublished
SSK 77	11-20 Feb 2015	Arabian Sea	Unpublished
SS379	19-23 December 2018	Arabian Sea	19
AS602D-615D	19-17 December 2007	Arabian Sea	55
SSD29	20-30 November 2016	Arabian Sea	Unpublished
SS379	6-18 December 2018	Arabian Sea	19
AS417-864	31 March to 10 April 2006	Bay of Bengal	14
AS802-864	27 Dec 2008-28 January 2009	Bay of Bengal	16
SSK24	15-22 March 2017	Bay of Bengal	Unpublished
SSK118	24 March to 1 April 2018	Bay of Bengal	62
SN142	1-20 June 2019	Bay of Bengal	Unpublished
Coastal fixed	September 2013 to December	Bay of Bengal	25
station	2017		
AS 402-AS416	19-30 March 2006	Bay of Bengal	14
SSD40	5-31 August 2017	Arabian Sea	34
SSD55	2-30 August 2018	Arabian Sea	34
ST001	1-9 <sup>th</sup> June 2020	Bay of Bengal	Unpublished

Table 1: The collection period, area and publication details of the measured aerosols composition samples used in this study.

Region	Region	Parameter	Winter	Spring	Summer	Fall	Annual
SWAS	53-58°E & 6-11°N	AAOD (/y)	0.003	0.004	-0.001 <sup>\$</sup>	0.005	0.002
		Mass (µg m <sup>-3</sup> y <sup>-1</sup> )	0.138	0.208	-0.062\$	0.260	0.103
		NO <sub>3</sub> <sup>-</sup> (µg m <sup>-3</sup> y <sup>-1</sup> )	0.004	0.006	-0.002\$	0.009	0.003
		$NH_4^+$ (µg m <sup>-3</sup> y <sup>-1</sup> )	0.007	0.011	-0.003\$	0.014	0.005
NWAS	58-61°E & 18-22°N	AAOD (/y)	0.002	0.002	-0.000*	0.006	0.002
		Mass (µg m <sup>-3</sup> y <sup>-1</sup> )	0.105	0.109	-0.009	0.338	0.109
		NO <sub>3</sub> <sup>-</sup> (µg m <sup>-3</sup> y <sup>-1</sup> )	0.003	0.003	-0.0003	0.010	0.003
		$NH_4^+$ (µg m <sup>-3</sup> y <sup>-1</sup> )	0.006	0.006	-0.001	0.018	0.006
NEAS	66-69°E & 18-22°N	AAOD (/y)	0.007	0.002	-0.000*	0.012	0.004
		Mass (µg m <sup>-3</sup> y <sup>-1</sup> )	0.364	0.115	-0.014	0.659	0.221
		NO <sub>3</sub> <sup>-</sup> (µg m <sup>-3</sup> y <sup>-1</sup> )	0.011	0.003	-0.0004	0.02	0.007
		$NH_4^+$ (µg m <sup>-3</sup> y <sup>-1</sup> )	0.019	0.006	-0.001	0.035	0.012
SEAS	73-76°E & 9-12°N	AAOD (/y)	0.013	0.013	0.0001*	0.012	0.008
		Mass (µg m <sup>-3</sup> y <sup>-1</sup> )	0.720	0.725	0.005*	0.685	0.443
		$NO_3^{-}(\mu g m^{-3} y^{-1})$	0.022	0.022	0.0002*	0.021	0.013
		$NH_4^+$ (µg m <sup>-3</sup> y <sup>-1</sup> )	0.038	0.038	0.0003*	0.036	0.023
NBoB	87-91°E & 17-21°N	AAOD (/y)	0.014	0.013	0.005	0.011	0.010
		Mass (µg m <sup>-3</sup> y <sup>-1</sup> )	1.667	1.523	0.593	1.33	1.176
		NO <sub>3</sub> <sup>-</sup> (µg m <sup>-3</sup> y <sup>-1</sup> )	0.049	0.045	0.018	0.039	0.034
		$NH_{4^{+}}(\mu g m^{-3} y^{-1})$	0.110	0.098	0.039	0.081	0.076
SEBoB	90-94°E	AAOD (/y)	-0.000*	0.001*	-0.000*	$0.002^{*}$	$0.000^{*}$
	& 4-8°N	Mass (µg m <sup>-3</sup> y <sup>-1</sup> )	-0.044	0.151	-0.053	0.246	0.049
		$NO_3^{-}(\mu g m^{-3} y^{-1})$	-0.001	0.004	-0.002	0.007	0.001
		$NH_4^+$ (µg m <sup>-3</sup> y <sup>-1</sup> )	-0.007*	0.012*	-0.001*	0.019\$	0.005*
SWBoB	82-86°E & 4-8°N	AAOD (/y)	0.002\$	0.002\$	-0.001 <sup>\$</sup>	0.003\$	0.001\$
		Mass (µg m <sup>-3</sup> y <sup>-1</sup> )	0.199	0.222	-0.067	0.346	0.346\$
		$NO_3^{-}(\mu g m^{-3} y^{-1})$	0.006	0.007	-0.002	0.010	0.010\$
		$NH_4^+$ (µg m <sup>-3</sup> y <sup>-1</sup> )	0.011\$	0.005*	0.002	0.031	0.013

Table 2: The slopes of long-term variability in AAOD, the concentration of mass, nitrate and ammonium in the aerosols at the selected regions in the northern Indian Ocean

\$ significant at p<0.05

\* not significant

All others were significant at p<0.001

Table 3: The external source of nitrogen and its contribution to new production in the Arabian Sea and Bay of Bengal

	1						
External	Arabian Sea			Bay of Bengal			Reference
C		0 (	ND				
Source	TgN/y	%	NP	TgN/y	%	NP	
			TgC/y			TgC/y	
N <sub>2</sub> fixation	$4.4\pm4$	70	29.0±26	$0.2\pm0.1$	7	$1.3\pm0.7$	55, 69, 70, 73-76
_							· · ·
Rivers	$0.1\pm0.03$	2	$0.7\pm0.2$	$1.7\pm0.4$	60	$11.2\pm3$	67
		_					
SGD	0.06+0.03	1	04+02	$0.1 \pm 0.02$	3	0.7+0.1	51
502	0.0020.00	-	0.120.2	0.120.02	2	0.7 =0.1	
Atm Den	17+04	27	11 2+3	0.9+0.2	30	6.0+1	This study
rum Dep	1.7_0.1	27	11.2±3	0.7_0.2	50	0.0±1	init study
Total	6 3+4		41 3+26	20+04		19 5+3	
1 Otul	0.5-4		+1. <u>J</u> ±20	2.0±0. <del>+</del>		17.5±5	

% is percentage contribution to total sources

SGD is Submarine Groundwater Discharge

NP is new production computed based on a C/N ratio of 6.6.

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### Legend to figures

Write NEAS, SEAS, NWAS, SWAS, NBoB, SWBoB and SEBoB to be given



## Figure 1: Monthly variations in anthropogenic AOD (AAOD) in the northern Indian Ocean

Figure 2: Distribution of measured a) mass, b) nitrate and c) ammonium concentrations during different periods in the northern Indian Ocean



Figure 3: Linear relationship of AAOD with a) mass, b) Nitrate and c) ammonium in the aerosols over the Arabian Sea (red) and Bay of Bengal (blue).



Figure 4: The estimated climatological monthly mean (2001-2020) mass of aerosols in the northern Indian Ocean



Figure 5: The estimated climatological monthly mean (2001-2020) nitrate concentrations in the aerosols in the northern Indian Ocean



Figure 6: The estimated climatological monthly mean (2001-2020) ammonium concentrations in the aerosols in the northern Indian Ocean



Figure 7: The estimated climatological monthly mean (2001-2020) nitrate depositional flux in the northern Indian Ocean



Figure 8: The estimated climatological monthly mean (2001-2020) ammonium depositional flux in the northern Indian Ocean.



Figure 9: a) Measured primary production integrated to the upper 5 m in the northern Indian Ocean, b) estimated primary production due to deposition of nitrate, c) estimated primary production due to deposition of ammonium and d) the contribution of primary production due to deposition of nitrate and ammonium to total primary production in the upper 5 m of the water column.



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